

Experimental shift of diet and DIC stable carbon isotopes: influence on shell $\delta^{13}\text{C}$ values in the Manila clam *Ruditapes philippinarum*

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1 Experimental shift of diet and DIC stable carbon isotopes: influence
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5 Poulain C.^a, Lorrain A.^a, Mas R.^b, Gillikin D.P.^c, Dehairs F.^b, Robert R.^d, Paulet Y-M.^a
6

7 ^a LEMAR, UMR CNRS/UBO/IRD 6539, Institut Universitaire Européen de la Mer, Place
8 Nicolas Copernic, 29280 Plouzané, France

9 ^b Earth System Sciences Research Group and Dept. of Analytical and Environmental
10 Chemistry, Vrije Universiteit Brussel, Pleinlaan 2, 1050 Brussels, Belgium

11 ^c Department of Earth Science and Geography, Vassar College, Poughkeepsie, New York
12 12604, USA

13 ^d IFREMER, UMR Laboratoire de Physiologie des Invertébrés, Station Expérimentale
14 d'Argenton, Presqu'île du vivier, 29840 Argenton, France
15
16

17 Corresponding Author:

18 Céline Poulain

19 LEMAR, UMR CNRS/UBO/IRD 6539, Institut Universitaire Européen de la Mer,
20 Technopole Brest-Iroise, Place Nicolas Copernic, 29280 Plouzané, France

21 Email: celine.poulain@univ-brest.fr

22 Phone number : + 33 2 98 49 88 38

23 Fax number : + 33 2 98 49 86 45
24
25

26 **Abstract**

27 The influences of diet and seawater dissolved inorganic carbon (DIC) on the carbon isotope
28 composition of shell aragonite ($\delta^{13}\text{C}_{\text{shell}}$) in the Manila clam *Ruditapes philippinarum* reared
29 under laboratory conditions were investigated. Clams were exposed to two successive
30 negative carbon isotope shifts: a first shift in diet ($\delta^{13}\text{C}_{\text{phytoplankton}}$) and a second shift, 35 days
31 later, in DIC ($\delta^{13}\text{C}_{\text{DIC}}$). Both successive shifts induced a decrease in $\delta^{13}\text{C}_{\text{shell}}$. These results are
32 the first to experimentally confirm an incorporation of respired carbon derived from food and
33 carbon from DIC into shell carbonate of adult bivalves. Skeletal $\delta^{13}\text{C}$ responded to changes in
34 the $\delta^{13}\text{C}$ of both diet and DIC in less than 7 days. Consequently, proxies based on $\delta^{13}\text{C}_{\text{shell}}$
35 may be used with high temporal resolution. Using $\delta^{13}\text{C}_{\text{phytoplankton}}$ as a proxy for the carbon
36 isotope composition of respired carbon ($\delta^{13}\text{C}_{\text{R}}$) resulted in a rather constant percentage of
37 metabolic carbon (C_{M}) into the shell carbonate over time (close to 12 %). However, an
38 accurate estimation of $\delta^{13}\text{C}_{\text{R}}$ is required in order to precisely estimate the percentage of
39 metabolic carbon incorporated into the shell. Despite the significant incorporation of
40 metabolic carbon into shell carbonate, our experimental results revealed that $\delta^{13}\text{C}_{\text{shell}}$ was
41 highly correlated with $\delta^{13}\text{C}_{\text{DIC}}$ ($r^2 = 0.77$, $p < 0.0001$). Thus it seems that $\delta^{13}\text{C}_{\text{shell}}$ is a
42 promising proxy of large scale variations in $\delta^{13}\text{C}_{\text{DIC}}$ and therefore of salinity in estuarine
43 water.

44
45

46 **Keywords:** $\delta^{13}\text{C}_{\text{shell}}$, DIC, salinity, metabolic carbon, *Ruditapes philippinarum*, estuarine
47 water, proxy

48
49

51 **1 Introduction**

52 Bivalves record environmental information in the form of chemical or structural messages
53 archived in their calcareous exoskeleton during growth. For this reason, they are widely used
54 for paleo-climatic and paleo-oceanographic reconstructions (e.g. Davenport, 1938, Weidman
55 et al., 1994, Tripathi et al., 2001, Dutton et al., 2002, Schöne et al., 2003, Lazareth et al.,
56 2006, Ivany et al., 2008, Goewert and Surge, 2008, Wanamaker et al., 2008). Epstein et al.
57 (1953) showed that oxygen isotope composition of mollusk shells ($\delta^{18}\text{O}_{\text{shell}}$) reflect the
58 temperature and oxygen isotope composition of the water ($\delta^{18}\text{O}_w$) in which they were formed.
59 This allows the use of shell carbonate $\delta^{18}\text{O}$ as a temperature proxy at least in cases where
60 $\delta^{18}\text{O}_w$ remains constant and is known (e.g. Weidman et al., 1994, Chauvaud et al., 2005,
61 Andrus and Rich, 2008). However, $\delta^{18}\text{O}_w$ is often strongly dependant on salinity (Epstein and
62 Mayeda, 1953, Craig and Gordon, 1965, Mook and Tan, 1991, Dettman et al., 2004).
63 Therefore, the estimation of $\delta^{18}\text{O}_w$ is difficult without an independent salinity proxy in
64 environments displaying large salinity variations such as estuaries. In well flushed estuaries
65 with short residence times, stable carbon isotope composition of dissolved inorganic carbon
66 ($\delta^{13}\text{C}_{\text{DIC}}$) is well correlated with salinity year round, especially when salinity is higher than 25
67 (Mook, 1971, Surge et al., 2001, Fry, 2002, Gillikin et al., 2006). As carbon precipitating
68 during shell mineralization is known to originate mainly from dissolved inorganic carbon
69 (DIC) of seawater (Mook and Vogel, 1968, McConnaughey et al., 1997, Gillikin et al., 2006,
70 McConnaughey and Gillikin, 2008, Owen et al., 2008), the carbon isotope composition of
71 bivalve shells ($\delta^{13}\text{C}_{\text{shell}}$) living in estuaries may possibly be used for reconstructing past
72 variations of salinity. A paleo-salinity proxy would also be useful for correcting paleo-
73 temperature based on $\delta^{18}\text{O}$ of estuarine bivalve shells (see Gillikin et al., 2006).

74 The carbon isotope composition of bivalve shell carbonates is, however, not only affected by
75 $\delta^{13}\text{C}_{\text{DIC}}$, but also by the physiology of the bivalve (Dillaman and Ford, 1982, Tanaka et al.,
76 1986, McConnaughey et al., 1997, Lorrain et al., 2004, Gillikin et al., 2006, 2007, 2009,
77 McConnaughey and Gillikin, 2008). Previous studies have shown that isotopically light
78 metabolic carbon, derived from food, is incorporated into shell carbonate (e.g. Tanaka et al.,
79 1986, McConnaughey et al., 1997, Wanamaker et al., 2007, Owen et al., 2008, Gillikin et al.,
80 2006, 2007, 2009). However, the processes of metabolic carbon incorporation into the shell
81 remain poorly known and the estimation of metabolic carbon contribution to the shell appears
82 highly variable between studies. McConnaughey et al. (1997), Lorrain et al. (2004) and

83 Gillikin et al. (2006) estimated the percentage of metabolic carbon integrated into the shell to
84 be lower than 10 % in deep-sea mollusks, *Pecten maximus* and *Mytilus edulis*, whereas
85 Gillikin et al. (2007, 2009) found values between 25 and 40 % in marine *Mercenaria*
86 *mercenaria* and freshwater Unionid shells, respectively. Consequently, development of
87 environmental proxies based on $\delta^{13}\text{C}_{\text{shell}}$ requires a better understanding of the processes
88 involved in the incorporation of metabolic carbon into the bivalve shell.

89 In the present study we performed a laboratory experiment using the Manila clam, *Ruditapes*
90 *philippinarum* (Adams and Reeve, 1850). This bivalve species, which has an aragonitic shell,
91 was chosen for several reasons. First, it is an euryhaline bivalve living mainly at salinity
92 levels ranging from 16 to 36 (Nie, 1991), buried a few centimeters in sandy and muddy
93 sediments in intertidal to subtidal zones. Because of its importance for aquaculture and
94 fisheries this species is well studied regarding its biology and physiology (e.g. Gouletquer et
95 al., 1989, Kim et al., 2001, Richardson, 1987, Marin et al., 2003, Flye-Sainte-Marie et al.,
96 2007). Moreover, clams, and particularly *Ruditapes* genus, are frequently found in shell-
97 middens, and are therefore suitable candidates as archives of past coastal human settlements
98 (Dupont and Marchand, 2008).

99 During this laboratory experiment, clams were reared under controlled values of
100 $\delta^{13}\text{C}_{\text{phytoplankton}}$ which were more depleted than those encountered in the field. Moreover, after
101 35 days of experiment, three salinity conditions and therefore three $\delta^{13}\text{C}_{\text{DIC}}$ were tested. The
102 aims of this study are (1) to demonstrate the incorporation of the two carbon sources (DIC
103 from surrounding seawater and DIC originating from respiration) into adult bivalve shells,
104 which, to our knowledge, has never been demonstrated experimentally, (2) to study the
105 response time of these carbon incorporations into the shell, and (3) to investigate the potential
106 use of $\delta^{13}\text{C}_{\text{shell}}$ as a proxy of high $\delta^{13}\text{C}_{\text{DIC}}$ variations and therefore of salinity in estuarine
107 water. In parallel, this experiment was designed to assess the effect of salinity variations on
108 clam physiology (i.e. condition index, shell growth rate and incorporation of carbon into soft
109 tissues) which has potential impacts on shell mineralization processes.

110

111 **2 Material and methods**

112 **2.1 Biological material**

113 A total of 250 Manila clams (*R. philippinarum*; two to three-years old; average length 27 mm;
114 $1\sigma = 2\text{mm}$) were collected by hand at low tide in an estuary located in the Gulf of Morbihan

115 (Bay of Kerdréan, 47°37'N, 2°56'W; Brittany; France; semidiurnal tidal regime) on the 18th
116 of August 2008. Clams were transferred to French Research Institute for Exploitation of the
117 Sea (IFREMER) Argenton Shellfish Laboratory (North Finistère, France).

118

119 **2.2 Experimental conditions**

120 Clams were randomly divided into three batches of 80 individuals and each batch was placed
121 into a 25 L tank without sediment. Water within the tanks was homogenized with an aquarium
122 pump. Each tank was supplied with UV sterilized water from a 300 L buffer tank allowing
123 renewal rate of 25% h⁻¹ and complete water changes were made once a week. The 300 L
124 buffer tanks were filled every 2 to 3 days with 1 µm filtered, UV sterilized water at room
125 temperature. The experimental system is illustrated in Fig. 1.

126 Water flowing into the 25 L tanks was supplemented continuously during the entire
127 experiment with cultured microalgae (50% *Isochrysis affinis galbana* (Tahitian strain *T.iso*)
128 and 50% *Chaetoceros calcitrans*) depleted in ¹³C ($\delta^{13}\text{C} = -58\text{‰}$; $1\sigma = 4\text{‰}$). ¹³C depleted
129 microalgae were obtained by bubbling ¹³C-depleted industrial CO₂ into the culture medium
130 (see Paulet et al., 2006 for elaboration). Algae were cultured at a salinity of 35. The
131 experiment was performed over 64 days, during which temperature (20°C) and photoperiod
132 (12/12) were kept constant.

133

134 During the first 35 days, salinity of the three tanks was maintained constant at 35. From days
135 35 to 64, one tank was kept at 35, while salinity was modified to 28 and 20 in the other two
136 tanks. Salinities of 20 and 28 were obtained by mixing seawater and tap water within the 300
137 L buffer tanks, and salinity was checked and adjusted using a conductivity meter (WTW, LF
138 197-S). Salinity was recorded every five minutes in the three tanks using an autonomous data
139 logger (YSI-600 OMS) from day 7 (t₇) to the end of the experiment (t₆₄).

140

141 **2.3 Sampling**

142 **2.3.1 Biological sampling**

143 **Soft tissue sampling**

144 During the first 35 days, while salinity was similar in all three tanks, two clams per tank were
145 collected each week. After the salinity change, five clams were randomly collected from each
146 tank weekly. Adductor and posterior muscle, digestive gland (without purging the digestive

147 tract) and mantle (without siphon) were dissected from each individual and were frozen at -
148 20°C. In order to determine initial isotope composition of tissues at the beginning of the
149 experiment (t_0), five clams from the field were processed in the same way. In addition, five
150 clams per tank were collected at t_{35} and t_{64} , and total soft tissues from each individual were
151 frozen at -20°C. At the end of the experiment, tissues were freeze-dried (48h) and shells were
152 air-dried. Tissues and shell were weighed in order to estimate the condition index (CI),
153 following Lucas and Beninger, (1985) :

$$154 \text{ CI} = (\text{Soft Tissues Dry Weight} / \text{Shell Dry Weight}) \times 100.$$

155

156 **Calcein marking and shell sampling**

157 Two calcein markings were used to establish a temporal scale in the shells. Clams in each
158 salinity treatment were exposed to a 150 mg L⁻¹ calcein solution during 4 hours (Rowley and
159 Mackinnon, 1995, Thébault et al., 2005) at days 15 (t_{15}) and 35 (t_{35}). At the end of the
160 experiment (t_{64}), 30 clams from each salinity were sacrificed and their shells were used for
161 growth rate determination and calcium carbonate sampling.

162

163 **2.3.2 Food and DIC sampling**

164 In order to determine the carbon isotope composition of the diet during the experiment, 15 mL
165 of the algal mix was filtered on a precombusted Whatman GF/F filter every week. Following
166 Lorrain et al. (2003), filters were immediately dried for 12h at 60°C and stored until analysis.
167 When the clams were collected, water was sampled in the field and 200 mL was filtered on a
168 precombusted Whatman GF/F filter to determine the carbon isotope composition of
169 particulate organic matter ($\delta^{13}\text{C}_{\text{POM}}$) using the method described above. The suspended POM
170 pool is a mixture of different sources of carbon (e.g. phytoplankton, microphytobenthos,
171 resuspended sediment, terrestrial carbon, marine micro-algae detritus). However, $\delta^{13}\text{C}_{\text{POM}}$
172 may be used as a proxy for $\delta^{13}\text{C}_{\text{phytoplankton}}$ (see Gillikin et al., 2006). Following Gillikin and
173 Bouillon (2007), water from the field and water within the 25 L clam tanks were sampled at t_0
174 and weekly respectively using 12 mL glass vials and poisoned with 20 μL of saturated
175 mercuric chloride (HgCl_2) solution until analysis of the DIC isotope composition.

176

177 **2.4 Shell preparation and calcium carbonate sampling**

178 For the purposes of calcium carbonate sampling, 30 shells were collected from each tank at
179 day 64 (end of experiment). A 8×10 mm piece of the shell, including the new shell material

180 formed during the experiment, was cut from the shell edge along the major growth axis and
181 embedded in epoxy resin (Araldite 2020, DIL, France). A cross-section (800 μm thick) was
182 then cut using a low speed diamond saw and glued on a glass slide with epoxy resin. Thick
183 sections were then ground with wet sandpaper (1200, 2400 and 4000 μm grit size) and
184 polished with a suspended diamond solution (3 μm). Shell sections were observed under a
185 fluorescence microscope (OLYMPUS BX41) equipped with a 50 W high-pressure Hg lamp
186 and a calcein filter. Photographs were acquired using a Hamamatsu C4742-95 digital camera
187 fitted on the microscope. The *R. philippinarum* shell consists of two aragonitic layers, an
188 inner homogeneous layer and an outer prismatic layer. Four evident marks in the outer layer
189 allowed assigning calendar dates on the shell sections: a cleft presumably due to the shell
190 growth stop induced by clam's collection in the field, the two calcein marks and the ventral
191 margin which sets the date the experiment was ended. For the 30 clams per tank collected on
192 the final day, growth between the four growth checks was measured on the shell section
193 photograph using the software Visilog 6.6. Five of the 30 shells (per batch) were selected
194 based on age (two-year old clams were selected to avoid any effect of ontogeny) and on high
195 shell growth rate to obtain sufficient material for carbon isotopic analysis. Since the tidal
196 periodicity of shell growth increment formation at intertidal flats is not lost in *R.*
197 *philippinarum* even after several months in constant laboratory conditions without emersion
198 (Richardson, 1987), it was possible to assign a calendar date to each growth increment (with
199 a precision of 3 days due to the behavioral disturbance induced by handling). Calcium
200 carbonates samples (average weight = 70 μg) were drilled to a depth of 350 μm in the middle
201 of the prismatic layer using a MicroMill (Merchantek) fitted with a 300 μm diameter drill bit.
202 One carbonate sample was drilled just before the cleft induced by clam collection in order to
203 measure the field isotopic value (t_0), one sample was drilled between the cleft and the first
204 calcein mark, two samples were drilled between the first and the second calcein mark and
205 three samples were drilled after the second calcein mark (see Fig. 2). Aragonite powder
206 samples were stored in glass vials until analysis.

207

208 **2.5 Isotopic analyses**

209 Freeze-dried tissues were ground to a homogeneous fine powder. 500 μg of the powder was
210 packed into 5 \times 9 mm precombusted silver cups. Carbon stable isotope analysis was performed
211 at the Earth System Sciences laboratory (Vrije Universiteit Brussel) using an Elemental
212 Analyzer (Flash 1112 EA Thermo Finnigan) coupled via a CONFLO III to a Thermo Delta V

213 Plus isotope ratio mass spectrometer (IRMS). A total of 42 samples of the international
214 isotopic reference standard IAEA-CH-6 (-10.45 ± 0.030 ‰) were analyzed with the samples
215 and yielded a reproducibility (1σ) of 0.36 ‰ for $\delta^{13}\text{C}$.

216 Filters were decarbonated via a 4-h exposure to HCl fumes in a dessicator (Lorrain et al.,
217 2003) and were packed into 5×9 mm tin cups. Isotopic analysis was performed as described
218 for the tissues.

219 The $\delta^{13}\text{C}$ of the total DIC was analyzed according to the protocol of Gillikin and Bouillon,
220 (2007). Briefly, 9.5 mL of water sample were acidified with 200 μL of pure orthophosphoric
221 acid in a 12 mL helium-flushed headspace vial, followed by overnight equilibration and
222 injection of 500 μL headspace into the carrier gas stream of the continuous flow EA-IRMS
223 (same instrument used for tissues and diet). The Miyajima et al. (1995) algorithm was used to
224 correct for the partitioning of CO_2 between headspace and water phase and to calculate the
225 $\delta^{13}\text{C}_{\text{DIC}}$. Along with the headspace injections a CO_2 house standard was injected ($n = 28$, 1σ
226 $= 0.2$ ‰). Based on this and an average reproducibility of DIC sample measurements of 0.11
227 ‰, a precision of the $\delta^{13}\text{C}_{\text{DIC}}$ better than 0.2 ‰ (1σ) can be expected (also see Gillikin and
228 Bouillon, 2007).

229 All carbonate isotopic analyses were performed on a Finningan MAT 252 IRMS equipped
230 with a Kiel III automated sampling device at the University of Arizona, USA. Samples were
231 reacted with > 100 % orthophosphoric acid at 70°C . The standard deviation of replicate
232 carbonate standards was 0.08‰ (1σ). All carbon isotopic results are reported relative to
233 VPDB (Vienna Pee Dee Belemnite) by calibration to the NBS-19 reference standard ($\delta^{13}\text{C} =$
234 $+1.95$ ‰).

235

236 **2.6 Data analysis**

237 All statistical analyses were performed using the software Statgraphics. Homoscedasticity was
238 tested using Bartlett's test ($\alpha=0.05$). ANOVAs were performed to check: (1) the differences in
239 mean shell growth of clams for the three salinity conditions between t_{35} and t_{64} , and (2) the
240 differences in carbon isotope composition between the different salinity conditions for each
241 organ at each sampling date. The differences in condition index between the three salinity
242 conditions were tested by performing non-parametric Kruskal-Wallis test.

243

244 **3 Results**

245 **3.1 Animal growth and condition during the experiment**

246 All clams exhibited significant growth during the experiment with shell growth between the
247 1st calcein marking (t_{15}) and the shell edge (t_{64}) of 1986 μm ($1\sigma = 1037 \mu\text{m}$; $n = 90$; see Fig.
248 2). The two bright calcein marks and the cleft induced by clam collection from the field were
249 readily identifiable in all shells. The three salinity treatments tested between day 35 (2nd
250 calcein marking) and the end of the experiment (shell edge) induced no significant differences
251 on shell growth (ANOVA, $p = 0.32$) with a mean shell growth of 1402 μm ($1\sigma = 752 \mu\text{m}$; $n =$
252 30), 1368 μm ($1\sigma = 598 \mu\text{m}$; $n = 30$) and 1135 μm ($1\sigma = 663 \mu\text{m}$; $n = 30$) at salinities 35, 28
253 and 20, respectively.

254 The condition index was almost constant during the entire experiment (average value = 8.80;
255 $1\sigma = 1.85$). Between days 35 and 64, at each sampling date, there were no significant
256 differences in average values of condition index between the three salinities conditions
257 (Kruskal-Wallis, $p > 0.05$).

258

259 **3.2 Salinity and $\delta^{13}\text{C}_{\text{DIC}}$**

260 During the first 35 days, salinity of the three tanks was 35.42 ($1\sigma = 0.21$), 35.77 ($1\sigma = 0.11$)
261 and 35.36 ($1\sigma = 0.10$) respectively. From days 35 to 64, the first tank was maintained at 35.28
262 ($1\sigma = 0.16$), while salinity was changed to 28.02 ($1\sigma = 0.14$) and 20.76 ($1\sigma = 0.13$) in the two
263 other tanks.

264 The carbon isotopic DIC values in the experimental tanks ranged between -1.28 ‰ and -3.34
265 ‰ at salinity 35, between -4.97 ‰ and -5.90 ‰ at salinity 28 and between -8.33 ‰ and -9.52
266 ‰ at salinity 20 (Fig. 3A). There was a high positive correlation between $\delta^{13}\text{C}_{\text{DIC}}$ and salinity
267 ($\delta^{13}\text{C}_{\text{DIC}} = 0.424 \times \text{salinity} - 17.04$, $r^2 = 0.95$, $p < 0.0001$, $n = 16$) as shown in Fig. 3B.

268

269 **3.3 $\delta^{13}\text{C}_{\text{tissues}}$**

270 From days 35 to 64, there was no significant effect of salinity on $\delta^{13}\text{C}$ values for digestive
271 gland and muscle at each sample date (ANOVA, $p > 0.05$). Therefore, values were pooled per
272 date as shown in Fig. 4. The $\delta^{13}\text{C}$ trend was similar for the different tissues (muscle, digestive
273 gland, mantle and total soft tissue) with a decrease over time. However, $\delta^{13}\text{C}_{\text{muscle}}$ decreased
274 from -14.7 to -32.6 ‰ in 64 days whereas $\delta^{13}\text{C}_{\text{digestive gland}}$ decreased from -17.5 to -38.5 ‰ in

275 only 7 days. Mantle $\delta^{13}\text{C}$ values were not significantly different than the muscle $\delta^{13}\text{C}$ value
276 (ANOVA, $p > 0.05$) at t_{35} . Total soft tissues $\delta^{13}\text{C}$ values were between muscle and digestive
277 gland $\delta^{13}\text{C}$ values.

278

279 **3.4 $\delta^{13}\text{C}_{\text{shell}}$**

280 Shell carbonate sampled before the cleft in the shell, which represents three to four days of
281 growth in the field just before collection, had mean $\delta^{13}\text{C}_{\text{shell}}$ values equal to -0.83‰ ($1\sigma =$
282 0.26‰) as shown in Fig. 5. At day 7, shell material formed in experimental tanks showed a
283 decreased in $\delta^{13}\text{C}_{\text{shell}}$ values in comparison to the field, with a mean value of -6.48‰ ($1\sigma =$
284 0.80‰ ; i.e., a 5.65‰ decrease on average). After the salinity changes, mean $\delta^{13}\text{C}_{\text{shell}}$ values
285 ranged from -5.03 to -6.64‰ for clams at salinity 35, from -8.35 to -9.82‰ for clams at
286 salinity 28 and from -10.39 to -12.11‰ for clams at salinity 20 (Fig. 5). These shell $\delta^{13}\text{C}$
287 values were significantly different between the three salinities at t_{40} , t_{52} and t_{62} (ANOVA, $p <$
288 0.0001). There was a significant positive linear relationship between $\delta^{13}\text{C}_{\text{shell}}$ and $\delta^{13}\text{C}_{\text{DIC}}$ ($r^2 =$
289 0.77 , $p < 0.0001$, $n = 83$) as shown in Fig. 6.

290

291 **4 Discussion**

292 At the beginning of the experiment, clams were transferred from the field where the $\delta^{13}\text{C}_{\text{POM}}$
293 value was -18‰ to laboratory conditions where $\delta^{13}\text{C}_{\text{phytoplankton}}$ was -58‰ . This shift in $\delta^{13}\text{C}$
294 of the clams diet led to a decrease in $\delta^{13}\text{C}$ in all tissues analyzed (muscle, digestive gland,
295 mantle and total soft tissues). This indicates a successful and quick incorporation of the
296 isotopically light dietary carbon from phytoplankton into clam tissues. In agreement with the
297 observation of Paulet et al. (2006) on *Pecten maximus* and *Crassostrea gigas* and Malet et al.
298 (2007) on *Crassostrea gigas*, our results indicate that the incorporation of dietary carbon is
299 faster in the digestive gland than in the muscle. This illustrates that the tissue turnover rate is
300 higher in the digestive gland than in the muscle.

301 In our study, there was no significant effect of salinity on the clams physiology as indicated
302 by the incorporation of carbon into the tissues, the condition index and shell growth rates.
303 These observations are in agreement with the results of Kim et al. (2001), who showed that
304 reasonable changes in salinity (35 to 20) have a small effect on activity rhythm and oxygen
305 consumption. Kim et al. (2001) report that Manila clams cannot maintain a normal metabolic

306 activity at salinities lower than 15 and all clams died after an exposure at a salinity of 5 for 7
307 days.

308

309 Our results showed that the transfer from field to experimental conditions was followed by a
310 decrease from -0.83 to -6.48 ‰ in $\delta^{13}\text{C}_{\text{shell}}$ in just 7 days (Fig. 5). Since the difference
311 between field and laboratory $\delta^{13}\text{C}_{\text{DIC}}$ was less than 1 ‰, the change in $\delta^{13}\text{C}_{\text{shell}}$ was most likely
312 due to the diet $\delta^{13}\text{C}$ shift. This strongly suggests an incorporation of carbon derived from food
313 into the shell. After 35 days, three salinity conditions were tested (20, 28 and 35). As
314 expected, shifts in salinity resulted in shifts in $\delta^{13}\text{C}_{\text{DIC}}$ (as we mixed waters with different
315 $\delta^{13}\text{C}_{\text{DIC}}$ values), which was followed by significant changes of $\delta^{13}\text{C}_{\text{shell}}$ values. These results
316 experimentally confirmed an integration into the shell of both carbon from seawater DIC and
317 metabolic DIC derived from food as suggested by number of authors (e.g. Tanaka et al., 1986,
318 McConnaughey et al., 1997, Lorrain et al., 2004, Wanamaker et al., 2007, McConnaughey
319 and Gillikin, 2008, Owen et al., 2008, Gillikin et al., 2006, 2007, 2009).

320

321 For both successive shifts ($\delta^{13}\text{C}_{\text{phytoplankton}}$ shift at t_0 and $\delta^{13}\text{C}_{\text{DIC}}$ shift at t_{35}), the incorporation
322 of carbon into the shell was very rapid: the first carbonate samples following both switches,
323 which correspond to an average of nearly 7 days, had significantly different $\delta^{13}\text{C}_{\text{shell}}$ than
324 carbonate samples before the shifts. Because seawater DIC is presumed to be supplied in
325 extrapallial fluids directly from the surrounding seawater through mantle epithelium and
326 periostracum (see McConnaughey and Gillikin, 2008), a short incorporation time is expected.
327 Interestingly the incorporation of carbon from food was also very rapid, which is not
328 necessarily expected because assimilation and metabolic processes may occur at a longer time
329 scale. This result suggests that during this experiment most respired C used in shell
330 construction is derived from the direct oxidation of food and is not from stored C sources.

331

332 During this experiment, $\delta^{13}\text{C}_{\text{DIC}}$ was not as constant as expected over time for the three
333 salinities (Fig. 3A). As salinity was constant during this experiment (except the voluntary shift
334 at t_{35}), $\delta^{13}\text{C}_{\text{DIC}}$ fluctuations are not explained by salinity variations. Two other hypotheses may
335 be considered. (1) Clam respiration and calcification will modify the CO_2 fluxes and then
336 $\delta^{13}\text{C}_{\text{DIC}}$, however, aquaria water was continuously renewed at a rate of $25\% \text{ h}^{-1}$ (see Materials
337 and methods), which would greatly reduce or even eliminate this effect on $\delta^{13}\text{C}_{\text{DIC}}$.
338 Nevertheless, future studies should consider this and analyze pH, total DIC, pCO_2 and even

339 O₂ levels as all these parameters have been proposed to influence shell $\delta^{13}\text{C}$ (see
340 McConnaughey and Gillikin, 2008 for a review); (2) Variable quantities of food supplied to
341 the clams is a more probable cause of these $\delta^{13}\text{C}_{\text{DIC}}$ fluctuations. In order to obtain the
342 appropriate phytoplankton concentration within the clam tanks, the flow of algae culture
343 supplied was adjusted daily as a function of algae culture concentration and the number of
344 clams in the tank. The $\delta^{13}\text{C}_{\text{DIC}}$ value in this algae culture was -22 ‰ during this experiment.
345 Therefore, when variable quantities of algae solutions are supplied to the clams,
346 corresponding variable ^{13}C -depleted DIC in the culture medium might induce differential
347 $\delta^{13}\text{C}_{\text{DIC}}$ values within the clam tanks. Despite the relative instability of $\delta^{13}\text{C}_{\text{DIC}}$, $\delta^{13}\text{C}_{\text{shell}}$
348 followed the same variations as $\delta^{13}\text{C}_{\text{DIC}}$. There was a positive linear relationship between
349 $\delta^{13}\text{C}_{\text{DIC}}$ and $\delta^{13}\text{C}_{\text{shell}}$ during the experiment with a high correlation coefficient value ($r^2 = 0.77$,
350 $p < 0.0001$, $n = 83$, Fig. 6). It seems that, in this species, $\delta^{13}\text{C}_{\text{shell}}$ may be a promising proxy of
351 large scale variations in $\delta^{13}\text{C}_{\text{DIC}}$ and therefore of salinity in estuarine water.

352

353 In order to estimate the percentage of metabolic carbon in the shell carbonate (C_M), previous
354 studies used the classical two end-member mixing equation proposed by McConnaughey et al.
355 (1997): $C_M = 100 \times (\delta^{13}\text{C}_{\text{shell}} - \epsilon_{\text{ar-b}} - \delta^{13}\text{C}_{\text{DIC}}) / (\delta^{13}\text{C}_R - \delta^{13}\text{C}_{\text{DIC}})$ where $\epsilon_{\text{ar-b}}$ is the enrichment
356 factor between bicarbonate and aragonite (+ 2.7 ‰; Romanek et al., 1992) and $\delta^{13}\text{C}_R$ is the
357 carbon isotope composition of respired carbon which has never been measured in bivalves.
358 According to Tanaka et al. (1986) and McConnaughey et al. (1997), $\delta^{13}\text{C}_{\text{tissues}}$ values provide
359 a good approximation of $\delta^{13}\text{C}_R$. Nevertheless, there is no consensus on the tissue used to
360 estimate $\delta^{13}\text{C}_R$; consequently we performed this estimation with all potential $\delta^{13}\text{C}_R$ values
361 ($\delta^{13}\text{C}$ of muscle, digestive gland, mantle, total soft tissues and cultured phytoplankton). There
362 was no significant difference between C_M values calculated for the three salinity conditions
363 with all potential $\delta^{13}\text{C}_R$ values and at any sampling date (Kruskal-Wallis, $p > 0.05$). The
364 similarity between C_M values at the three salinities seems to reject a strong impact of carbonic
365 anhydrase (CA) activity on shell $\delta^{13}\text{C}$. Low salinity could result in decreased carbonic
366 anhydrase (CA) activity in the animal, as has been described for some bivalves (Henry and
367 Saintsing, 1983). In addition, CA activity has also been shown to be inhibited by Cl^- ions
368 (Pocker and Tanaka, 1978) and therefore by salinity levels. Gillikin et al. (2006) hypothesized
369 that CA, which catalyses the reaction of bicarbonate to CO_2 , thereby facilitating diffusion of
370 DIC through membranes (Paneth and O'Leary, 1985), may add or remove carbon species
371 from the internal calcifying fluids. A reduction in CA activity due to low salinity could cause
372 a reduction in environmental DIC entering the animal, resulting in a larger relative

373 contribution of metabolic DIC and more negative $\delta^{13}\text{C}$ in the calcifying fluid and shell. Since
374 there was no significant difference between C_M values calculated for the three salinity
375 conditions, it appears that the salinity effect on CA activity and the ability of the animal to
376 move environmental DIC into their calcification space is small or non-existent. The similarity
377 between C_M values at the three salinities also seems to reject a strong impact of possible
378 varying O_2 and CO_2 levels on shell $\delta^{13}\text{C}$ (see McConnaughey and Gillikin, 2008) that could
379 arise with changing salinities. However, future studies should measure these parameters to
380 estimate their role.

381 The estimated values of C_M shown in Fig. 7 were highly sensitive to the type of tissue used
382 for calculation and, at t_7 , C_M values ranged from 21 % using digestive gland to 50 % using
383 muscle (note that we calculated C_M for each date based on $\delta^{13}\text{C}_{\text{tissues}}$ values pooled from all
384 salinities as there was no salinity effect on C_M). The estimations using $\delta^{13}\text{C}_{\text{mantle}}$ and $\delta^{13}\text{C}_{\text{total}}$
385 *soft tissues* resulted in intermediate values that ranged between the estimations made using
386 muscle and digestive gland $\delta^{13}\text{C}$ values. At the end of this experiment, these differences were
387 reduced (12 % using digestive gland to 20 % using muscle). Interestingly C_M values
388 calculated with digestive gland or muscle, as a proxy for respired carbon $\delta^{13}\text{C}$, between the
389 clams collected at the end of the experiment and clams collected in the field, did not greatly
390 differ (see Fig. 7, 17 vs 12 % or 21 vs 20 % using the digestive gland and muscle
391 respectively). As clams in the laboratory were growing without sediments, this would suggest
392 that ^{13}C depleted porewater plays little to no role in biomineralization (see McConnaughey
393 and Gillikin, 2008). Indeed, incorporation of ^{13}C depleted porewater in the field would lead to
394 more negative shell $\delta^{13}\text{C}$ and then an overestimation of C_M in the field. In our lab experiment,
395 without such depleted porewaters, C_M values should have been much smaller than the field if
396 carbon from these waters was strongly incorporated into the shells.

397 The C_M fluctuations through time were higher using $\delta^{13}\text{C}_{\text{muscle}}$ as $\delta^{13}\text{C}_R$ than using $\delta^{13}\text{C}_{\text{digestive}}$
398 *gland* (50 to 20 % using muscle; 21 to 12 % using digestive gland). Interestingly, the estimation
399 using $\delta^{13}\text{C}_{\text{phytoplankton}}$ (which is lower than $\delta^{13}\text{C}_{\text{tissues}}$, especially at the beginning of the
400 experiment) results in an even more constant estimation of C_M over time than using $\delta^{13}\text{C}_{\text{digestive}}$
401 *gland*, with a value close to 12 %. Since salinity variations didn't affect the clams physiology
402 and other experimental conditions such as temperature and food quantity were constant during
403 this experiment, such a constant value can be expected. Moreover, a C_M value of 12 % is near
404 the typical value of 10 % recorded for most aquatic mollusks (see McConnaughey and
405 Gillikin, 2008). Assuming that $\delta^{13}\text{C}_R$ is close to $\delta^{13}\text{C}_{\text{phytoplankton}}$ during this experiment
406 illustrates that metabolic CO_2 originates from very reactive tissues or directly from food

407 oxidation when animals are fed continuously, and that muscle tissues were not appropriate to
408 estimate C_M .

409 In the field, availability of energy (food) and energy allocation within bivalves is highly
410 variable throughout the year (Paulet et al., 2006). During phytoplankton blooms, when
411 metabolism is generally high, more of the carbon used for shell mineralization may be derived
412 from metabolic CO_2 which mainly results from oxidation of food. In such a case, $\delta^{13}C_R$ may
413 be closer to $\delta^{13}C$ values of phytoplankton. During less productive periods, carbon contribution
414 may mainly come from DIC and because metabolism is mainly fuelled from reserves, $\delta^{13}C_R$
415 should be closer to $\delta^{13}C$ values of storage organs (e.g. muscle). Thus, in bivalves, both
416 proportion and carbon isotope value of metabolic carbon may be variable over the year. If this
417 hypothesis is true, then studies investigating high frequency variations in $\delta^{13}C_{shell}$ should use
418 different tissues to estimate $\delta^{13}C_R$ according to the time period studied (digestive gland or
419 phytoplankton during productive periods and muscle during less productive period). Studies
420 that average more time (e.g. annual) can use tissues with a slower turnover rate, such as the
421 muscle. However, under natural conditions the effect would be much smaller than in this
422 study as in the field food items are not likely to vary as much as the 40 ‰ shift imposed in
423 this study, and muscle and digestive gland $\delta^{13}C$ values would not be so isotopically different
424 (e.g. Lorrain et al., 2002).

425

426 Our study highlights that $\delta^{13}C$ values of respired carbon are necessary to accurately determine
427 the percentage of metabolic carbon integrated into bivalve shells. This respired CO_2 is
428 however very difficult to sample and has never been measured in bivalves.

429

430

431 **Conclusion**

432 This study experimentally highlights the incorporation of two carbon sources into shell
433 carbonate: dissolved inorganic carbon from seawater and respired carbon derived from food.
434 The incorporation time of both carbon sources into the shell is very short (less than 7 days).
435 These short incorporation times indicate that proxies based on $\delta^{13}C_{shell}$ may be used with high
436 temporal resolution. Our experimental results also demonstrate that (1) the physiology of this
437 euryhaline clam was not affected by the salinity treatments ranging between 20 and 35 during
438 this experiment; (2) salinity does not seem to affect carbon cycling within the clam; and (3)
439 $\delta^{13}C_{shell}$ is highly correlated to $\delta^{13}C_{DIC}$. Consequently, $\delta^{13}C_{shell}$ of *R. philippinarum* is a

440 promising proxy of large scale variations in $\delta^{13}\text{C}_{\text{DIC}}$ and therefore of salinity in estuarine
441 water.

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444

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460 **References**

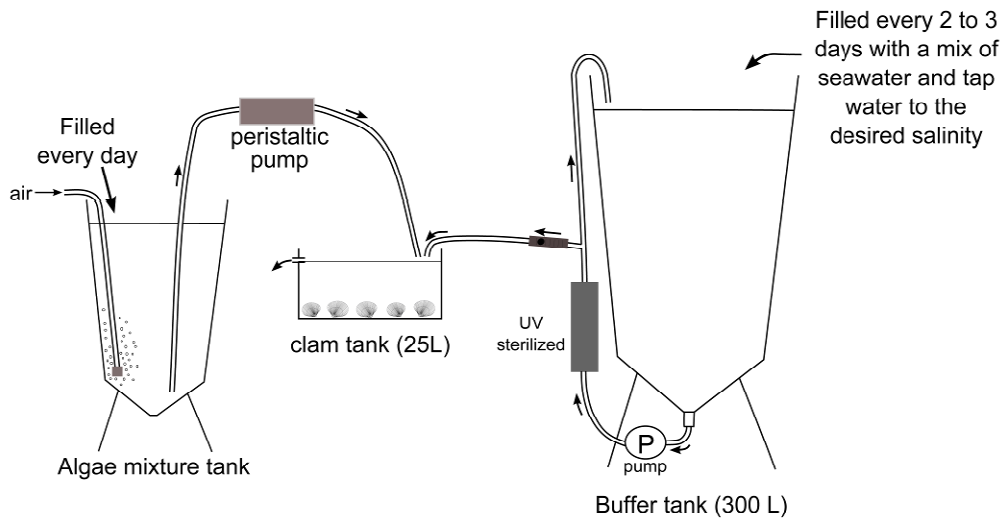
- 461 Andrus, C.F.T. and Rich, K.W., 2008. A preliminary assessment of oxygen isotope
462 fractionation and growth increment periodicity in the estuarine clam *Rangia cuneata*.
463 *Geo-Mar. Lett* 28 (5-6), 301-308.
- 464 Chauvaud, L., Lorrain, A., Dunbar, R.B., Paulet, Y.-M., Thouzeau, G., Jean, F., Guarini, J.-
465 M. and Mucciarone, D., 2005. Shell of the Great Scallop *Pecten maximus* as a high-
466 frequency archive of paleoenvironmental changes. *Geochem. Geophys. Geosyst.* 6
467 (Q08001), doi:10.1029/2004GC000890.
- 468 Craig, H. and Gordon, L.I., 1965. Deuterium and oxygen 18 variations in the ocean and
469 marine atmosphere. In: E. Tongiorgi (Editor), *Stable Isotopes in Oceanographic
470 Studies and Paleotemperatures* Consiglio Nazionale delle Ricerche, Laboratorio di
471 *Geologia Nucleare*, Pisa, pp. 9-130.
- 472 Davenport, C., 1938. Growth lines in fossil pectens indicators of past climates. *J. Paleontol.*
473 12 (5), 514-515.
- 474 Dettman, D.L., Flessa, K.W., Roopnarine, P.D., Schone, B.R. and Goodwin, D.H., 2004. The
475 use of oxygen isotopic variation in shells of estuarine mollusks as a quantitative record
476 of seasonal and annual Colorado River discharge. *Geochim. Cosmochim. Acta* 68 (6),
477 1253-1263.
- 478 Dillaman, R.M. and Ford, S.E., 1982. Measurement of calcium-carbonate deposition in
479 mollusks by controlled etching of radioactively labeled shells. *Mar. Biol.* 66, 133-143.

- 480 Dupont, C. and Marchand, G., 2008. Coastal exploitation on the neolithic of Western France:
481 La pointe Saint-Gildas (Préfailles). *Environmental archeology* 13 (2), 143-152.
- 482 Dutton, A.L., Lohmann, K.C. and Zinsmeister, W.J., 2002. Stable isotope and minor element
483 proxies for Eocene climate of Seymour Island, Antarctica. *Paleoceanography* 17 (2),
484 doi: 10.1029/2000PA000593.
- 485 Epstein, S., Buchsbaum, R., Lowenstam, H.A. and Urey, H.C., 1953. Revised carbonate-water
486 isotopic temperature scale. *Geol. Soc. Am. Bull.* 64, 1315-1326.
- 487 Epstein, S. and Mayeda, T., 1953. Variations of ^{18}O content of waters from natural sources.
488 *Geochim. Cosmochim. Acta* 4, 213-224.
- 489 Flye-Sainte-Marie, J., Jean, F., Paillard, C., Ford, S., Powell, E., Hofmann, E. and Klinck, J.,
490 2007. Ecophysiological dynamic model of individual growth of *Ruditapes*
491 *philippinarum*. *Aquaculture* 266, 130-143.
- 492 Fry, B., 2002. Conservative mixing of stable isotopes across estuarine salinity gradients: a
493 conceptual framework for monitoring watershed influences on downstream fisheries
494 production. *Estuaries Coasts* 25 (2), 264-271.
- 495 Gillikin, D.P. and Bouillon, S., 2007. Determination of $\delta^{18}\text{O}$ of water and $\delta^{13}\text{C}$ of dissolved
496 inorganic carbon using a simple modification of an elemental analyzer – isotope ratio
497 mass spectrometer (EA-IRMS): an evaluation. *Rapid Commun. Mass Spectrom.* 21,
498 1475-1478.
- 499 Gillikin, D.P., Hutchinson, K.A. and Kumai, Y., 2009. Ontogenic increase of metabolic
500 carbon in freshwater mussel shells. *J. Geophys. Res. -Biogeosciences* 114 (G01007),
501 doi:10.1029/2008JG000829.
- 502 Gillikin, D.P., Lorrain, A., Bouillon, S., Willenz, P. and Dehairs, F., 2006. Stable carbon
503 isotopic composition of *Mytilus edulis* shells: relation to metabolism, salinity, $\delta^{13}\text{C}_{\text{DIC}}$
504 and phytoplankton. *Org. Geochem.* 37, 1371-1382.
- 505 Gillikin, D.P., Lorrain, A., Meng, L. and Dehairs, F., 2007. A large metabolic carbon
506 contribution to the $\delta^{13}\text{C}$ record in marine aragonitic bivalve shells. *Geochim.*
507 *Cosmochim. Acta* 71, 2936-2946.
- 508 Goewert, A. and Surge, D., 2008. Seasonality and growth patterns using isotope
509 sclerochronology in shells of the Pliocene scallop *Chesapecten madisonius*. *Geo-Mar.*
510 *Lett.* 28, 327-338.
- 511 Gouletquer, P., Heral, M., Deslous-Paoli, J.M., Prou, J., Garnier, J., Razet, D. and
512 Boromthanarat, W., 1989. Ecophysiology and energy balance of the Japanese clam
513 *Ruditapes philippinarum*. *J. Exp. Mar. Biol. Ecol.* 132, 85-108.
- 514 Henry, R.P. and Saintsing, D.G., 1983. Carbonic-anhydrase activity and ion regulation in 3
515 species of osmoregulating bivalve mollusks. *Physiol. Zool.* 56 (2), 274-280.
- 516 Ivany, L.C., Lohmann, K.C., Hasiuk, F., Blake, D.B., Glass, A., Aronson, R.B. and Moody,
517 R.M., 2008. Eocene climate record of a high southern latitude continental shelf:
518 Seymour Island, Antarctica. *Geol. Soc. Am. Bull.* 120, 659-678.
- 519 Kim, W.S., Huh, H.T., Huh, S.H. and Lee, T.W., 2001. Effects of salinity on endogenous
520 rhythm of the Manila clam, *Ruditapes philippinarum* (Bivalvia: Veneridae). *Mar. Biol.*
521 138 (1), 157-162.
- 522 Lazareth, C.E., Lasne, G. and Ortlieb, L., 2006. Growth anomalies in *Protothaca thaca*
523 (Mollusca, Veneridae) shells as markers of ENSO conditions. *Clim. Res.* 30 (3), 263-
524 269.
- 525 Lorrain, A., Paulet, Y.-M., Chauvaud, L., Dunbar, R.B., Mucciarone, D. and Fontugne, M.,
526 2004. $\delta^{13}\text{C}$ variation in scallop shells: Increasing metabolic carbon contribution with
527 body size? *Geochim. Cosmochim. Acta* 68 (17), 3509-3519.

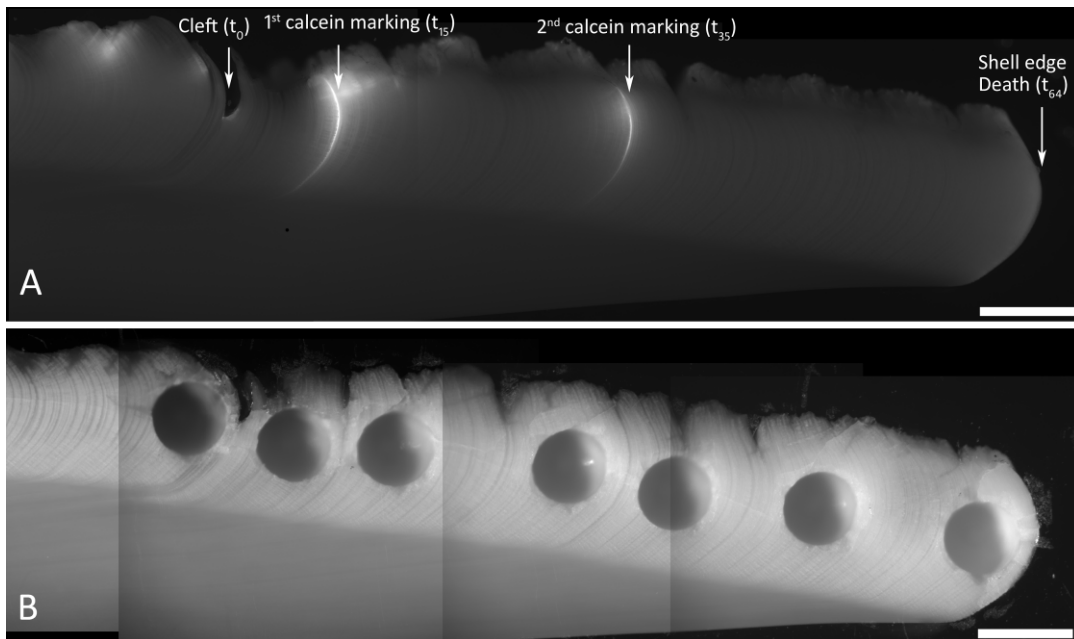
- 528 Lorrain, A., Paulet, Y.-M., Chauvaud, L., Savoye, N., Donval, A. and Saout, C., 2002.
529 Differential $\delta^{13}\text{C}$ and $\delta^{15}\text{N}$ signatures among scallop tissues: implications for ecology
530 and physiology. *J. Exp. Mar. Biol. Ecol.* 275 (1), 47-61.
- 531 Lorrain, A., Savoye, N., Chauvaud, L., Paulet, Y.M. and Naulet, N., 2003. Decarbonation and
532 preservation method for the analysis of organic C and N contents and stable isotope
533 ratios of low-carbonated suspended particulate material. *Anal. Chim. Acta* 491 (2),
534 125-133.
- 535 Lucas, A. and Beninger, P.G., 1985. The use of physiological indices in marine bivalve
536 aquaculture. *Aquaculture* 44, 187-200.
- 537 Malet, N., Sauriau, P.-G., Faury, N., Soletchnik, P. and Guillou, G., 2007. Effect of seasonal
538 variation in trophic conditions and the gametogenic cycle on $\delta^{13}\text{C}$ and $\delta^{15}\text{N}$ levels of
539 diploid and triploid Pacific oysters *Crassostrea gigas*. *Mar. Ecol.-Prog. Ser.* 346, 203-
540 217.
- 541 Marin, M.G., Moschino, V., Deppieri, M. and Lucchetta, L., 2003. Variations in gross
542 biochemical composition, energy value and condition index of *T. philippinarum* from
543 the Lagoon of Venice. *Aquaculture* 219, 858-871.
- 544 McConnaughey, T.A., Burdett, J., Whelan, J.F. and Paull, C.K., 1997. Carbon isotopes in
545 biological carbonates: respiration and photosynthesis. *Geochim. Cosmochim. Acta* 61
546 (3), 611-622.
- 547 McConnaughey, T.A. and Gillikin, D.P., 2008. Carbon isotopes in mollusk shell carbonates.
548 *Geo-Mar. Lett.* 28, 287-299.
- 549 Miyajima, T., Yamada, Y., Hanba, Y.T., Yoshii, K., Koitabashi, T. and Wada, E., 1995.
550 Determining the stable isotope ratio of total dissolved inorganic carbon in lake water
551 by GC/C/IRMS. *Limnol. Oceanogr.* 40 (5), 994-1000.
- 552 Mook, W.G., 1971. Paleotemperature and chlorinities from stable carbon and oxygen isotopes
553 in shell carbonate. *Palaeogeogr. Palaeoclimatol. Palaeoecol.* 9, 245-263.
- 554 Mook, W.G. and Tan, F.C., 1991. Stable carbon isotopes in rivers and estuaries. In: E.T.
555 Degens, S. Kempe and J.E. Richey (Editors), *Biogeochemistry of Major World Rivers*.
556 John Wiley and Sons Ltd, pp. 245-264.
- 557 Mook, W.G. and Vogel, J.C., 1968. Isotopic equilibrium between shells and their
558 environment. *Science* 159 (3817), 874-875.
- 559 Nie, Z.Q., 1991. The culture of marine bivalve mollusks in China. In: W. Menzel (Editor),
560 *Estuarine and marine bivalve mollusk culture*. CRC Press, Boston, pp. 261-276.
- 561 Owen, E.F., Wanamaker, A.D., Feindel, S.C., Schöne, B.R. and Rawson, P.D., 2008. Stable
562 carbon and oxygen isotope fractionation in bivalve (*Placopecten magellanicus*) larval
563 aragonite. *Geochim. Cosmochim. Acta* 72, 4687-4698.
- 564 Paneth, P. and O'Leary, M.H., 1985. Carbon isotope effect on dehydration of bicarbonate ion
565 catalyzed by carbonic-anhydrase. *Biochemistry* 24, 5143-5147.
- 566 Paulet, Y.M., Lorrain, A., Richard, J. and Pouvreau, S., 2006. Experimental shift in diet $\delta^{13}\text{C}$:
567 A potential tool for ecophysiological studies in marine bivalves. *Org. Geochem.* 37
568 (10), 1359-1370.
- 569 Pocker, Y. and Tanaka, N., 1978. Inhibition of carbonic-anhydrase by anions in carbon
570 dioxide-bicarbonate system. *Science* 199, 907-909.
- 571 Richardson, C.A., 1987. Tidal bands in the shell of the clam *Tapes philippinarum*. *Proc. R.*
572 *Soc. Lond. B* 230, 367-387.
- 573 Romanek, C.S., Grossman, E.L. and Morse, J.W., 1992. Carbon isotopic fractionation in
574 synthetic aragonite and calcite: effects of temperature and precipitation rate. *Geochim.*
575 *Cosmochim. Acta* 56 (1), 419-430.

- 576 Rowley, R.J. and Mackinnon, D.I., 1995. Use of the fluorescent marker calcein in
577 biomineralisation studies of brachiopods and other marine organisms. Bulletin de
578 l'Institut Oceanographique, Monaco 14, 111-120.
- 579 Schöne, B.R., Oschmann, W., Rössler, J., Castro, A.D.F., Houk, S.D., Kröncke, I., Dreyer,
580 W., Janssen, R., Rumohr, H. and Dunca, E., 2003. North Atlantic Oscillation
581 dynamics recorded in shells of a long-lived bivalve mollusk. *Geology* 31 (12), 1037-
582 1040.
- 583 Surge, D., Lohmann, K.C. and Dettman, D.L., 2001. Controls on isotopic chemistry of the
584 American oyster, *Crassostrea virginica*: implication for growth patterns. *Palaeogeogr.*
585 *Palaeoclimatol. Palaeoecol.* 172, 283-296.
- 586 Tanaka, N., Monaghan, M.C. and Rye, D.M., 1986. Contribution of metabolic carbon to
587 mollusk and barnacle shell carbonate. *Nature* 320, 520-523.
- 588 Thébault, J., Chauvaud, L., Clavier, J., Fichez, R. and Morize, E., 2005. Evidence of a 2-day
589 periodicity of striae formation in the tropical scallop *Comptopallium radula* using
590 calcein marking. *Mar. Biol.* 149 (2), 257-267.
- 591 Tripathi, A., Zachos, J.C., Marinovich, L. and Bice, K., 2001. Late Paleocene Arctic coastal
592 climate inferred from molluscan stable and radiogenic isotope ratios. *Palaeogeogr.*
593 *Palaeoclimatol. Palaeoecol.* 70, 101-113.
- 594 Wanamaker, A.D., Kreutz, K.J., Borns, H.W., Introne, D.S., Feindel, S., Funder, S., Rawson,
595 P.D. and Barber, B.J., 2007. Experimental determination of salinity, temperature,
596 growth, and metabolic effects on shell isotope chemistry of *Mytilus edulis* collected
597 from Maine and Greenland. *Paleoceanography* 22 (2).
- 598 Wanamaker, A.D., Kreutz, K.J., Schöne, B.R., Pettigrew, N., Borns, H.W., Introne, D.S.,
599 Belknap, D., Maasch, K.A. and Feindel, S., 2008. Coupled North Atlantic slope water
600 forcing on Gulf of Maine temperature over the past millennium. *Clim. Dyn.* 31 (2),
601 183-194.
- 602 Weidman, C.R., Jones, G.A. and Lohmann, K.C., 1994. The long-lived mollusk *Artica*
603 *islandica* - a new paleoceanographic tool for the reconstruction of bottom
604 temperature for the continental shelves of the northern North-Atlantic Ocean. *J.*
605 *Geophys. Res.* 99 (C9), 18305-18314.
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609 **Figures**
610

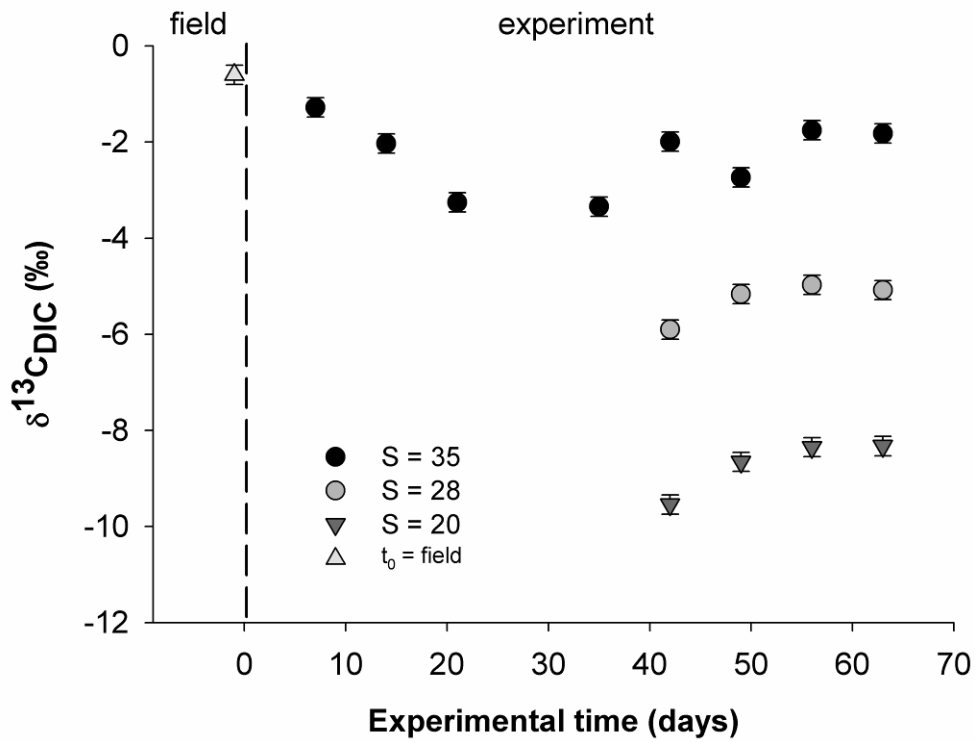


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612 Fig. 1. Schematic diagram of the experimental system. The complete experimental system
613 consisted of three clam and buffer tanks, one for each clam batch. The same algae culture was
614 used for the three clam batches.
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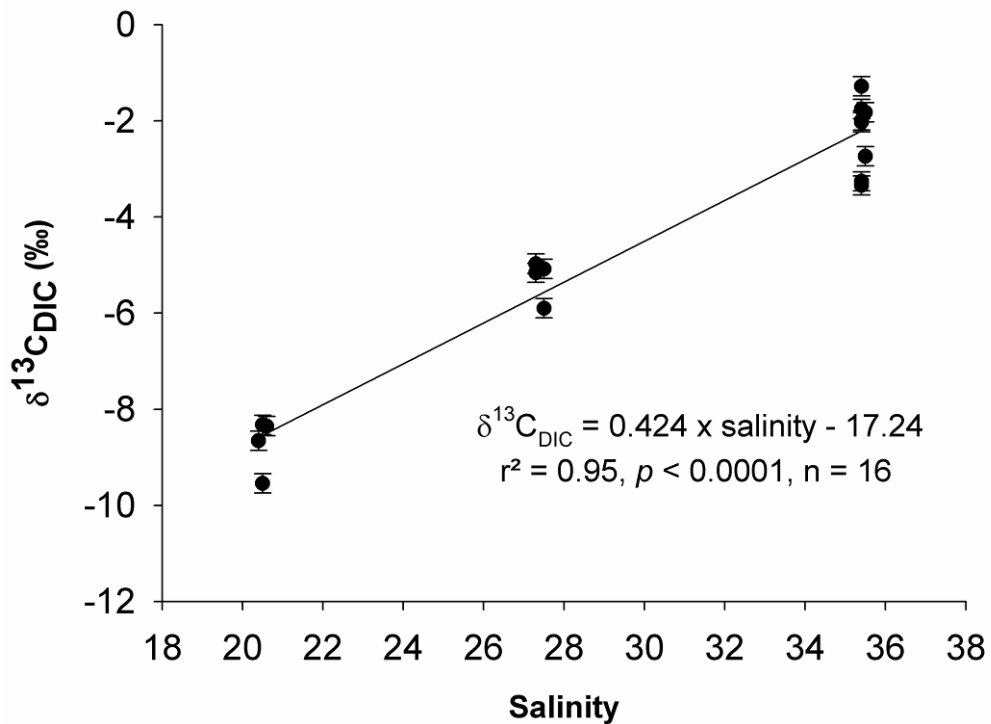


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617 Fig. 2. Photographs of shell transversal section. Growth direction is left to right, the outside of
618 the shell is toward the top of picture. A: The cleft corresponds to a growth stop due to clam
619 harvesting in the field on the 18th of August 2008. Clams were kept in laboratory for 64 days.
620 The two calcein stains at t_{15} and t_{35} are visible. Shell edge corresponds to the end of the
621 experiment. B: Drill holes (300 μm diameter; 350 μm depth) made in the middle of the outer
622 layer. Scale bar = 500 μm .

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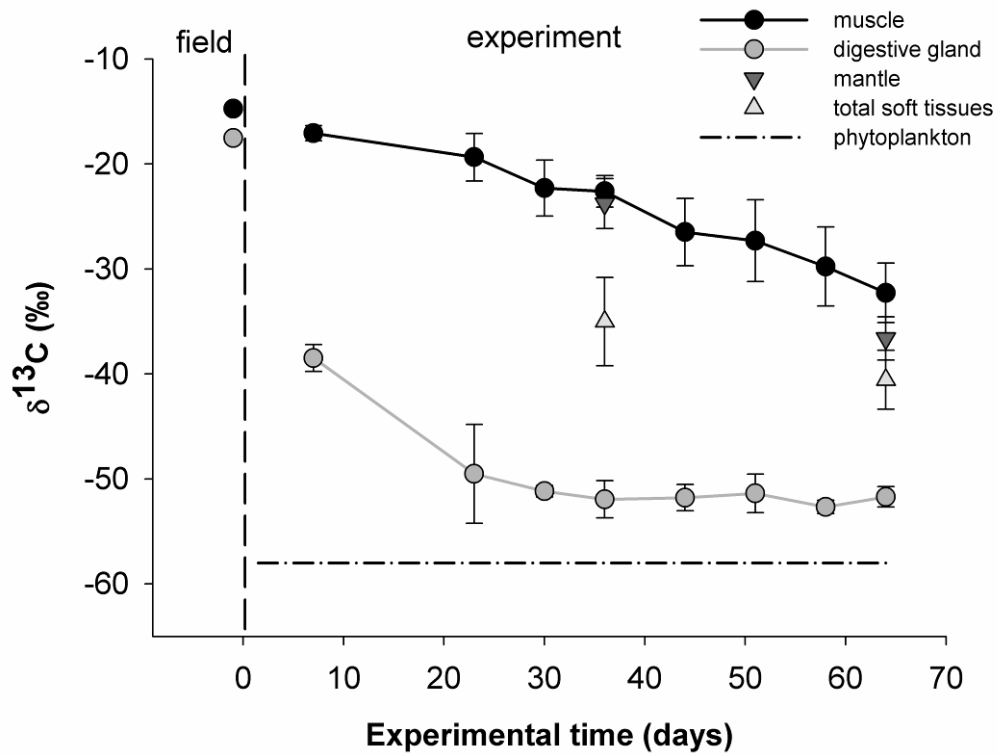
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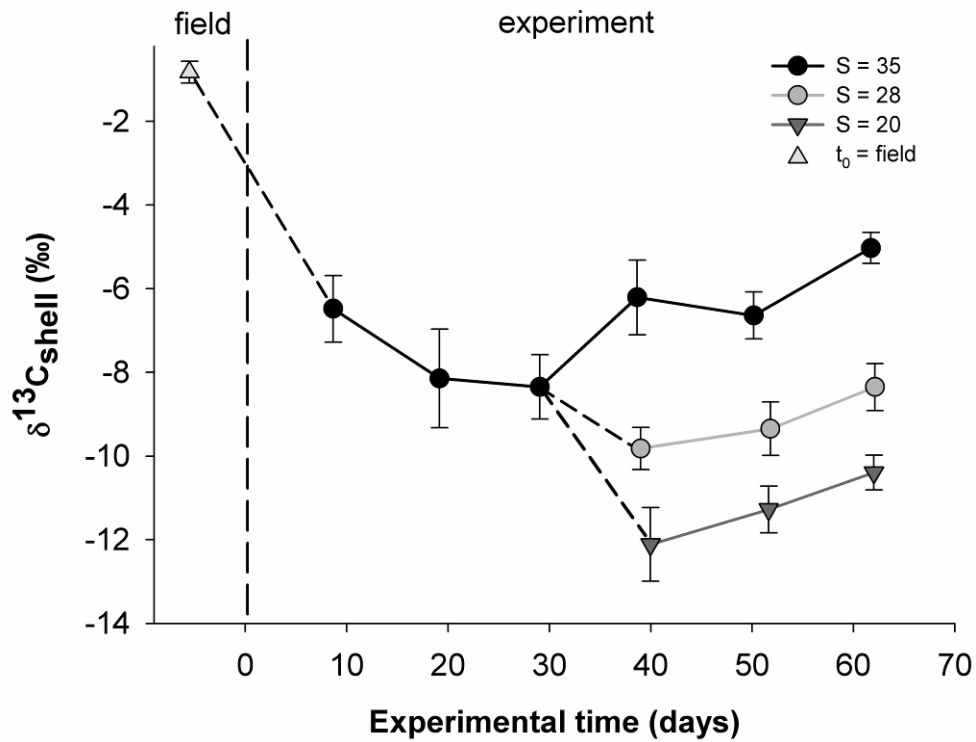
626 Fig. 3. A: DIC stable carbon isotope values ($\delta^{13}\text{C}_{\text{DIC}}$) during the experiment. Three salinity
627 conditions (20, 28 and 35) were tested from days 35 to 64. B: $\delta^{13}\text{C}_{\text{DIC}}$ as a function of salinity
628 during the experiment. Error bars represent the analytical precision (1 σ).

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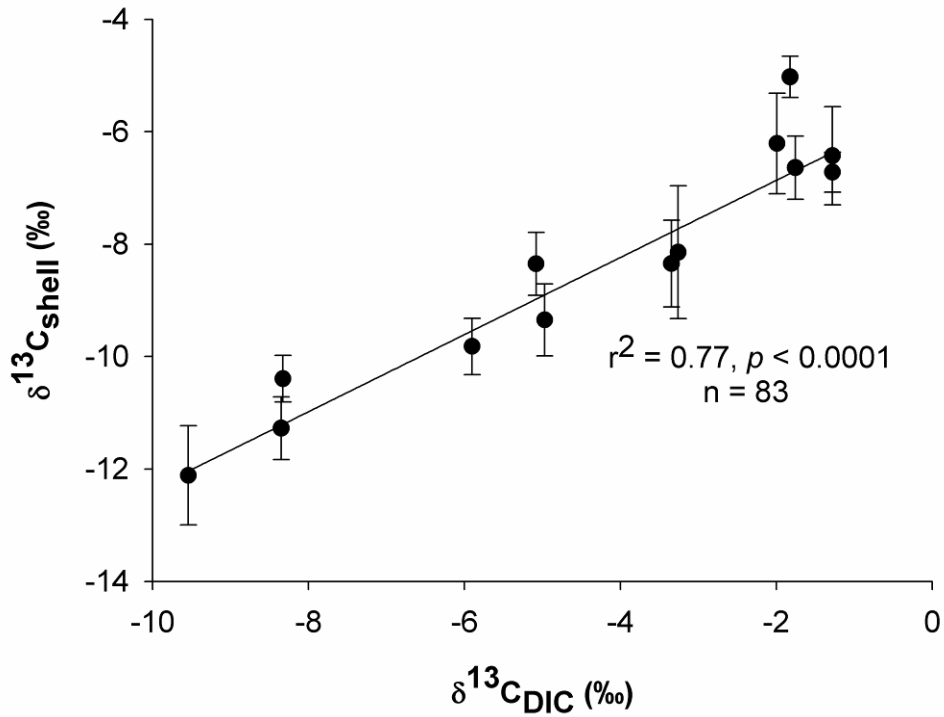
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Fig. 4. Evolution of stable carbon isotope values ($\delta^{13}\text{C}$) of Manila clam tissues (muscle, digestive gland, mantle and total soft tissues) and phytoplankton during the experiment. Values are expressed as means and error bars represent standard deviations, $n = 5$ during the first 35 days; $n = 15$ from day 35 to 64 because the results for the three different salinities were pooled.



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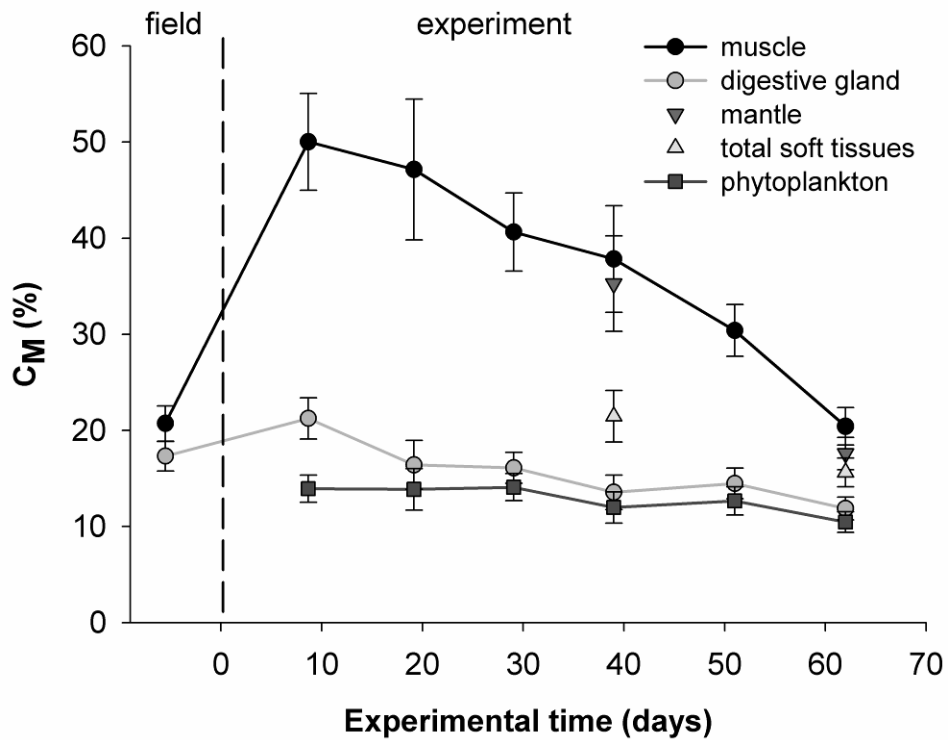
Fig. 5. Evolution of stable carbon isotope values ($\delta^{13}C_{shell}$) in *R. philippinarum* shells during the experiment for the three salinity conditions (20, 28 and 35). Values are expressed as means and error bars represent standard deviations, n = 15 during the first 35 days; n = 5 from day 35 to 64 for each point.



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645 Fig. 6. $\delta^{13}\text{C}_{\text{shell}}$ vs $\delta^{13}\text{C}_{\text{DIC}}$ in clam tanks for samples taken over the experiment. Error bars
 646 represents standard deviations. The solid line shows the linear least squares regression
 647 calculated with all data obtained during this experiment ($r^2 = 0.77$, $p < 0.0001$, $n = 83$).

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651 Fig. 7. Estimates of the proportions of metabolic carbon (C_M) incorporated into shell
 652 carbonates using the equation of McConnaughey et al. (1997). Muscle, digestive gland,
 653 mantle, total soft tissues and phytoplankton carbon isotope composition were used to
 654 approximate $\delta^{13}C_R$. Values are expressed as means and error bars represent standard
 655 deviations. Different patterns are observed depending on the tissue used for the approximation
 656 of $\delta^{13}C_R$.