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A seasonal study of diatom dynamics in the North Atlantic during the POMME experiment (2001): Evidence for Si limitation of the spring bloom

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[1] The Si cycle in the northeast Atlantic (40°–43.5°N, 15°–21°W) was investigated in the winter, spring, and late summer during the Programme Océan Multidisciplinaire Méso Echelle (POMME) (2001). The aim of this study was to determine the principal causes of the onset and subsequent decline of the diatom bloom, with an emphasis on nutritional limitation processes. The siliceous compartment dynamics was characterized through silicic acid distribution, size-fractionated biogenic silica and Si uptake rates, Si uptake kinetics (K_S and V_{max}), and export rates. The results of the POMME cruises indicated an important seasonal variability of the siliceous compartment in the North Atlantic together with a spatial variability following a south-north gradient in parallel to the increasing stock of nutrients. Here we present the first kinetic evidences for Si limitation of the diatom spring bloom in the northeast Atlantic, supporting previous reports of potentially limiting silicic acid concentrations in this region. Integrated Si uptake rates were very low throughout the survey, except at the northern anticyclonic eddy site, where a *Pseudo-nitzschia* bloom was observed during spring and ranged between 0.04 and 11.2 mmol m⁻² d⁻¹, which is comparable to values obtained in oligotrophic regions. The overall annual Si production budget for the entire POMME area (375,000 km²) was extrapolated to 9.8×10^9 mol Si yr⁻¹, while Si export fluxes at 400 m only represented 3% of the surface production.

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1. Introduction

[2] The intense spring bloom occurring each year in the North Atlantic is often depicted as the most important biological signal visible from space [Esaias *et al.*, 1986]. The combined effect of the CO₂ solubility pump and the biological pump confers to this oceanic area the status of a large CO₂ sink, which was estimated to range from 6 to 8×10^{12} gC yr⁻¹ [Takahashi *et al.*, 2002]. In the eastern North Atlantic, the winter mixed layer depth discontinuity zone divides the basin into two regions: north of this discontinuity, the bloom is thought to progress northward following the onset of the spring stratification as described in Sverdrup's critical depth model [Sverdrup, 1953]; south of this

boundary, the shallower winter mixing brings less nutrients to surface waters, making this more oligotrophic region more likely to be nutrient than light-limited [Siegel *et al.*, 2001]. The south-north density gradient is furthermore disrupted by a large mesoscale activity in this region, generating eddies generally less than 100 km wide, which induces a secondary spatial variability of nutrient stocks. Hence the northward progression of the spring bloom is not homogeneous but has rather been depicted as a moving patchwork of mesoscale features characterized by different bloom stages [Karrasch *et al.*, 1996].

[3] The North Atlantic spring bloom is usually dominated by diatoms supporting pulse-like sedimentation events [Lampitt, 1985; Berger and Wefer, 1990; Lampitt and Antia, 1997], which are thought to contribute for a large part to the export of organic carbon to the deep ocean [Smetacek, 1985; Honjo and Manganini, 1993]. Diatoms distinguish themselves from other phytoplanktonic groups by their absolute requirement for dissolved Si, necessary to their metabolism and cell wall formation [Sullivan and Volcani, 1981]. A lot of interest is usually given to the role of nitrogen in the progression of the spring bloom, while relatively fewer studies focus on the role of silicon, which is necessary to diatom growth. Silicic acid depletion has been reported as the

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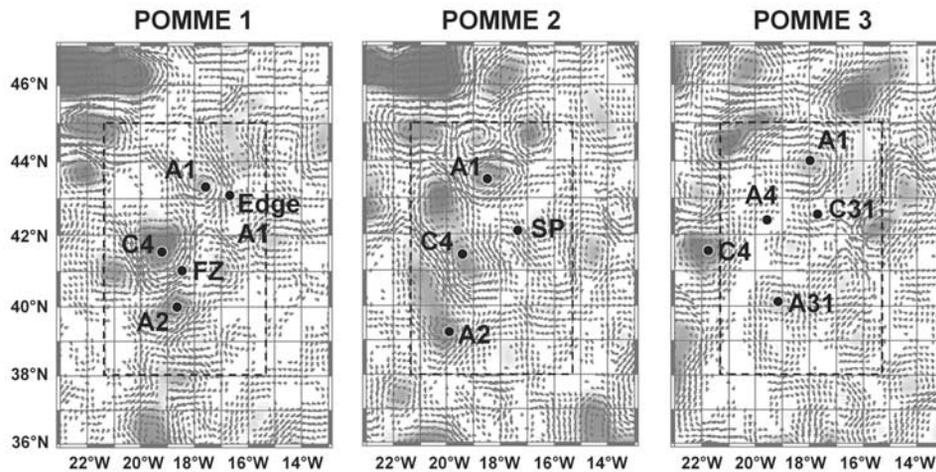


Figure 1. Synoptic current maps at 100 m for the second legs were derived from the SOPRANE model (M. Assenbaum, LEGOS/CNRS). Cyclonic and anticyclonic eddies were positioned and numbered (C, cyclonic; A, anticyclonic). The study sites (black dots) sampled during the second legs were repositioned according to the hydrological features.

principal cause for the decline of the spring bloom in several hydrological regimes including the North Atlantic [Lochte *et al.*, 1993; Sieracki *et al.*, 1993]. However, this oceanic region remains largely undersampled with regards to silicic acid distribution [Louanchi and Najjar, 2001] and estimates of Si uptake rates are scarce. In the North Atlantic, Si production rates were quantified in the western Atlantic mid-ocean gyres [Brzezinski and Nelson, 1989, 1996; Brzezinski and Kosman, 1997], in a coastal African upwelling [Nelson and Goering, 1978] and more recently in the northeast Atlantic between Iceland and Scotland [Brown *et al.*, 2003].

[4] In this article, we present a comprehensive study of the Si cycle dynamics in several mesoscale structures of the eastern Atlantic during a winter-summer transition survey. The Programme Océan Multidisciplinaire Méso Echelle (POMME) study area was located between the Azores Island and the Iberian peninsula, encompassing a region characterized by strong mesoscale activity as well as by a meridional gradient located at 41°–41.5°N, where an eastward current extended through the POMME domain [Giordani *et al.*, 2005; Mémerly *et al.*, 2005].

[5] The aim of this study was to determine the principal cause of the onset and subsequent decline of the diatom bloom, with an emphasis on nutritional limitation processes. The diatom bloom progression was investigated during winter, spring and late summer 2001 in the 38°–45°N, 15°–21°W area at several 1-D study sites. The siliceous compartment dynamics was characterized through silicic acid distribution, size-fractionated biogenic silica and Si uptake rates, Si uptake kinetics (K_S and V_{max}) and Si export rates. Here, we present the first kinetic evidences for Si limitation of the diatom spring bloom in the northeast Atlantic, supporting previous reports of potentially limiting silicic acid concentrations in this region.

2. Methods

2.1. Study Area and Sampling Strategy

[6] The POMME program was conducted in the northeast Atlantic on board R/V *L'Atalante* and *Thalassa* (IFREMER)

in 2001. The three POMME cruises (P1, P2, P3) were each divided into two legs covering the winter, spring and late summer seasons. The first legs were dedicated to an extensive coverage of the study area while the second legs were aimed at performing process studies within specific hydrological features. Mesoscale structures such as cyclonic or anticyclonic eddies, and frontal areas were resolved by combining data from floats and drifters deployed into the area (Marvors, Provors and Surdrifts), CTD and XBT data obtained during the first legs, as well as altimetric sea level measurements (see details of Mémerly *et al.* [2005]). Data assimilation and modeling using the SOPRANE model [Assenbaum and Reverdin, 2005] yielded synoptic current maps and allowed to accurately relocate each study site on its corresponding hydrological feature (Figure 1). This study presents the data acquired during the second legs for different mesoscale features (Table 1), some of which were transient while others could be followed throughout the year such as: a northern anticyclonic eddy identified as A1, a cyclonic eddy identified as C4 and a southern anticyclonic eddy identified as A2. This last structure unfortunately drifted out of the study area during P3 and was replaced by a smaller anticyclonic eddy at the same latitude, A31. A general drift of most eddies toward the southwest was noticeable between the cruises [Le Cann *et al.*, 2005].

[7] Each study site was sampled during 3 days. Arriving on site, a drifting line equipped with two sediment traps (200 and 400 m) was immersed for a 48 hour period. On the first day, particulate and dissolved silica were sampled with a rosette sampler on a deep CTD cast (0–1000 m). On the second day, the 0–100 m layer was sampled before dawn for the same parameters, and supplementary whole water samples for silicic acid uptake measurements were attached to a drifting line at the depth of collection and incubated in situ for 24 hours.

2.2. Dissolved and Particulate Matter

[8] Major nutrient measurements (NO_3^- , PO_4^{3-}) were performed at every station and prior to the tracer incubation

Table 1. Conductivity-Temperature-Depth (CTD) Casts Coordinates and Hydrological Characteristics of the Study Sites During the Second Legs of the Three Programme Océan Multidisciplinaire Méso Echelle (POMME) Cruises (P1, P2, P3)^a

Hydrological Structures	CTD	Longitude, deg	Latitude, deg	Date	Salinity	SST, °C	Z_{eup} , m
<i>P1: Winter</i>							
A2	1090	18°44	40°07	2 March 2001	35.86	14.4	58
FZ	1117	18°37	41°03	7 March 2001	35.85	14.1	55
C4	1146	19°14	41°46	11 March 2001	35.68	13.2	56
A1	1173	17°22	43°14	15 March 2001	35.70	13.0	47
Edge A1	1184	16°21	43°08	17 March 2001	35.66	12.7	45
<i>P2: Spring</i>							
A2	2091	19°45	39°45	19 April 2001	35.90	15.3	50
C4	2179	19°42	41°52	23 April 2001	35.74	14.4	45
SP	2273	17°34	42°01	27 April 2001	35.79	14.3	50
A1	2410	18°49	43°20	2 May 2001	35.76	14.0	48
<i>P3: Late Summer</i>							
A31	3094	19°20	40°09	20 Sept. 2001	36.24	21.7	86
A4	3163	19°51	42°13	24 Sept. 2001	36.03	20.8	93
C4	3247	21°57	41°28	28 Sept. 2001	35.89	20.8	91
C31	3331	17°58	42°26	8 Oct. 2001	35.87	17.9	88
A1	3339	17°33	44°19	3 Oct. 2001	35.80	19.1	69

^aThe mesoscale structures that were sampled are defined by letters (A, anticyclonic eddy; C, cyclonic eddy) and by numbers for each new eddy. Apart from eddies, a frontal zone (FZ) and an area of no eddy activity (SP, Saddle Point) were also sampled. SST, sea surface temperature; Z_{eup} , surface salinities and the depths of the euphotic zone.

experiments. Samples were taken before dawn in a 20 mL polyethylene flask. All samples were analyzed on board immediately after sampling according to *Tréguer and Le Corre* [1975], using a Technicon Autoanalyser[®]. The detection limits are 0.05 μM and 0.02 μM for nitrate and phosphate respectively. Silicic acid samples were preserved with 25 μL mercuric chloride (HgCl_2) and stored at 4°C until analysis at the laboratory. Dissolved silicon was measured manually on a spectrophotometer following *Brzezinski and Nelson* [1986] on a CECIL spectrofluorometer, with a detection limit of 0.05 μM .

[9] Pigment samples were filtered onto 0.7 μm filters (Whatman GF/F, 25 mm) and stored in liquid nitrogen until analysis. Samples were extracted in 3 mL HPLC-grade methanol and injected onto a reversed phase C8 Hypersil MOS column (dimensions: 3 \times 100 mm, 3 μm pore size) and analyzed with an Agilent Technologies 1100 series HPLC system coupled with a Thermoquest AS3000 autosampler according to the protocol described by *Claustre et al.* [2004]. Precision of the method averaged 5% and detection limits for TChl a and fucoxanthin (a pigment found in diatoms) were above 0.001 mg m^{-3} .

[10] Particulate biogenic silica (BSi) samples were filtered on cascading filter towers onto 0.6 μm and 10 μm filters (Nucleopore, 47 mm) allowing size fractionation. Filters were dried at room temperature on board and assayed for biogenic silica following the hot NaOH digestion technique described by *Nelson et al.* [1989]. BSi trap samples were measured following the same method and the final Si flux was corrected for the initial input of silicic acid in the cup with the formaldehyde solution and for BSi dissolution after collection within each cup. Immediately after sample filtration for BSi assays, the concentration of dissolved silicic acid remaining in the filtrate was measured. These values were corrected for the initial amount of dissolved Si present in the filtered seawater used to fill the sampling cups and

then added to the particulate BSi values to correct for dissolution of diatom frustules after collection.

2.3. Si Uptake Rate Profiles and Kinetic Parameters

[11] Dawn-to-dawn (24 hours) in situ uptake experiments were realized using an immersed production line at eight incubation depths chosen over the 0–100 m layer. Whole seawater samples (275 mL) were collected before dawn and spiked with 0.022 μCi (800 Bq) of the radioactive isotope silicon-32 (^{32}Si), which specific activity was 15.6 $\text{kBq } \mu\text{g Si}^{-1}$. After incubation, samples were filtered on cascading filter towers onto 0.6 μm and 10 μm filters (Nucleopore, 47 mm) allowing size fractionation. Filters were rinsed with filtered (0.2 μm) seawater, and placed in scintillation vials. The ^{32}Si uptake was measured in a Packard 1600-TR scintillation counter by Cerenkov effect, following the method described by *Tréguer et al.* [1991] and *Leynaert* [1993]. Precision of the method averages 10%. Si specific uptake rates V_{Si} (in d^{-1}) have been derived from BSi (in $\mu\text{mol L}^{-1}$) and ρSi (absolute Si uptake rates in $\mu\text{mol L}^{-1} \text{d}^{-1}$) measurements using the following equation:

$$V_{Si} = \frac{\rho Si}{BSi}. \quad (1)$$

[12] Kinetic uptake experiments were carried out at 20 and 40 m at all sites except during P3. The water samples were collected from the same Niskin bottles as those used for in situ incubation. Samples were spiked with ^{32}Si and received increasing Na_2SiO_3 additions ranging from 0.5 to 50 μM . Bottles were incubated on deck for 8 hours in flow-through incubators maintained at sea surface temperature and covered with nickel screens (22.5 and 8% of incident irradiance). Samples were thereafter treated as described for in situ samples. Kinetic parameters V_{max} and K_S were calculated by fitting the data to a Michaelis-Menten curve using Sigmaplot[®].

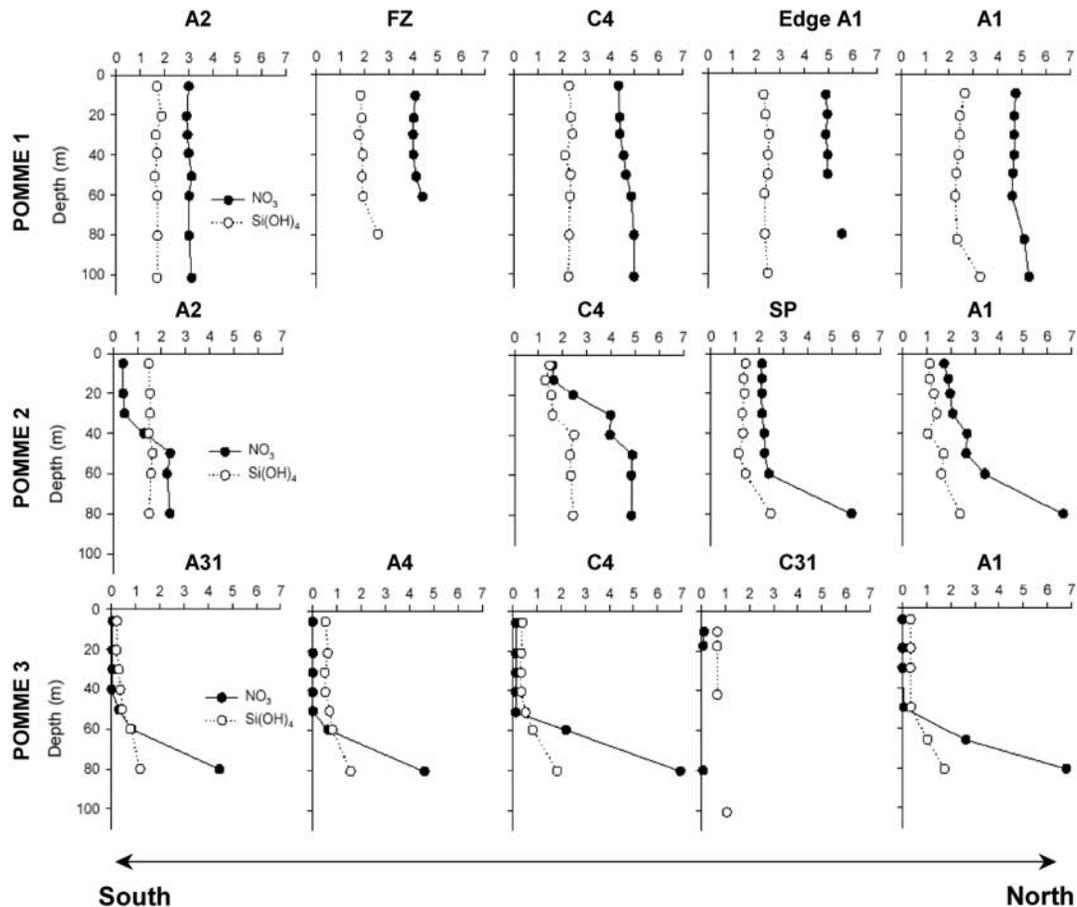


Figure 2. Vertical silicic acid and nitrate profiles (μM) for the 0–100 m layer during the three cruises POMME 1, 2, and 3.

2.4. Primary Production

[13] Primary production (PP) rates and integrated primary production (ΣPP) were measured in situ using the ^{14}C method according to *Moutin and Raimbault* [2002]. No PAR (Photosynthetically Active Radiation) profiles were available for the POMME cruises, but the depth of the euphotic zone was derived from the integrated Chl *a* content of the water column according to *Morel and Berthon* [1989].

3. Results

3.1. Hydrological Features

[14] The water mass composed between the surface and the permanent deep thermocline (500–800 m) was entirely constituted of the Eastern North Atlantic Central Water (ENACW), identified by salinities >35.5 . The ENACW are also characterized by temperatures $>10^\circ\text{C}$ and rising to values ranging from 13° to 22°C in the late summer [*van Aken*, 2001]. The study area was characterized by the presence of several mesoscale cyclonic and anticyclonic eddies as well as by an eastward current located at 41° – 41.5°N . This current, characterized by an intense isopycnal displacement, was a quasi-permanent structure at least until June 2001, but presented at individual times large meanders around various eddy structures [*Mémery et al.*, 2005].

[15] A full hydrographic description of these structures is given in companion papers [*Le Cann et al.*, 2005; *Giordani et al.*, 2005; *Reverdin et al.*, 2005; *Fernández et al.*, 2005] and is beyond the scope of this paper. For each study site, the sea surface temperatures (SST), the surface salinities and the depths of the euphotic zone (Z_{eu}) are reported in Table 1. In March, surface temperatures ranged from 12.7° to 14.4°C , increased by approximately 1°C during spring season and increased further by approximately 6°C in the late summer. Z_{eu} values were intermediate during winter (45–58 m), decreased slightly during spring bloom (45–50 m) except at the anticyclone A1 site where it remained equivalent, and increased during the oligotrophic late summer period (69–91 m).

3.2. Nutrient Stocks

[16] Nitrate and silicic acid profiles in the 0–100 m layer for each site are presented in Figure 2. In the winter (P1), silicic acid concentrations ranged from 1.6 to $2.5 \mu\text{M}$ and were always lower than nitrate concentrations, with an average difference of 1.5 – $2 \mu\text{M}$. Both surface nitrate and silicic acid availability were lowest in the south and increased progressively to the north. In the spring (P2), nutrients decreased significantly but were not completely depleted. Nitrate concentrations remained below $2 \mu\text{M}$ and silicic acid concentrations ranged from 1 to $1.5 \mu\text{M}$. Silicic acid levels

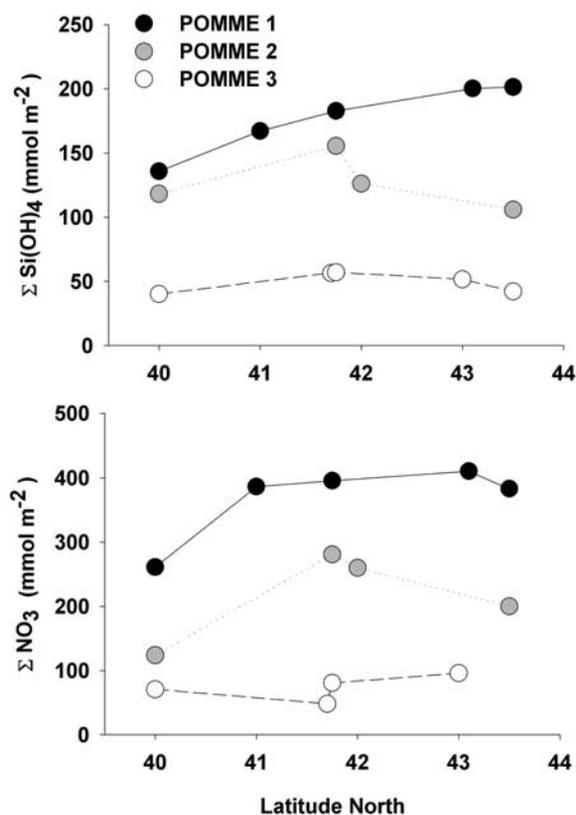


Figure 3. Integrated silicic acid and nitrate stocks along the south-north gradient (40°–44°N) in mmol m^{-2} .

were again lower than nitrate except at anticyclone A2, where nitrate was on average $0.5 \mu\text{M}$ in the first 30 m. In the late summer (P3), nutrients were severely depleted over the first 50–60 m over the entire study area. In the first 50 m, nitrate was inferior to the detection limit ($0.05 \mu\text{M}$) at all sites and silicic acid ranged from $0.2 \mu\text{M}$ in A2 to $0.5 \mu\text{M}$ in A1.

[17] Integrated nutrient stocks (over 80 m) were determined at all sites and compared for each cruise in order to assess the seasonal evolution of biological drawdown (Figure 3). During winter, the integrated silicic acid stock increased by 30% from south to north, with the highest stock in A1 ($200 \text{ mmol Si m}^{-2}$). During spring, less than 15% of the winter silicic acid stock was depleted in the southern part of the study area (A2 and C4) while Si utilization increased up to 47% in the northern region (A1). In the late summer, Si stocks were low and homogeneous over the entire region ($40\text{--}57 \text{ mmol m}^{-2}$). Integrated nitrate stocks followed approximately the same distribution, with a south-north increasing gradient during winter and a stronger depletion in A1 than in C4 during spring. However, the rate of depletion was much higher between P1 and P2 than for silicic acid, with a 50% decrease of the nitrate stock at both A2 and A1, and a 30% decrease in C4. In the late summer, nitrate was depleted in the surface layer over the entire region and stocks ranged from 48 mmol N m^{-2} in C4 to 96 mmol N m^{-2} in A1.

3.3. Integrated Particulate Matter Distribution and Primary Production

[18] Total biogenic silica (BSi) profiles are presented for the three mesoscale structures which were revisited

throughout the year (A1, C4 and A2) for the 0–100 m layer (Figure 4). The winter BSi concentrations were low and homogeneous over the 0–60 m layer ($0.04\text{--}0.14 \mu\text{mol L}^{-1}$), but slightly higher than the late summer values ($0.01\text{--}0.06 \mu\text{mol L}^{-1}$), indicating that diatoms were already growing by the end of February. By April–May (P2), the diatom spring bloom was well developed, with a peak BSi concentrations of $0.61 \mu\text{mol L}^{-1}$ at 10 m in C4 and a deeper maximum of $0.88 \mu\text{mol L}^{-1}$ between 30 and 60 m in A1. In A2 however, BSi concentrations remained much lower, reflecting the lower silicic acid stock utilization, and a low maximum value of $0.18 \mu\text{mol L}^{-1}$ was observed at 40 m. Size-fractionated integrated BSi ($0.6\text{--}10 \mu\text{m}$ and $>10 \mu\text{m}$) was measured at each site during winter and spring, but were not carried out in the late summer due to the very low diatom abundance (Figure 4). In the winter, the contribution of small diatoms ($<10 \mu\text{m}$) increased regularly from south to north, representing 34% in A2, 48% in C4 and 69% in A1. During spring, large diatoms $>10 \mu\text{m}$ were slightly dominant in both A2 and C4 (with an equal contribution of 58%) and constituted up to 76% of the total BSi stock in A1.

[19] Total integrated BSi, pigment stocks and primary production (ΣPP) for the three revisited sites are presented in Figure 5. A strong seasonal pattern was observed for both BSi and fucoxanthin stocks, with a strong increase between P1 and P2. BSi stocks presented a weak spatial variability during winter, with fairly similar concentrations ranging from 4 to 8 mmol Si m^{-2} . On the other hand, fucoxanthin stocks showed an increase from south to north already in the winter, as stocks more than tripled in A1 compared to A2. During spring, the intersite variability rose sharply and BSi stocks increased by a factor 2, 3, and 7 from south to north respectively. The overall highest stock during the POMME cruises was measured in A1 in May ($61 \text{ mmol Si m}^{-2}$). This latitudinal gradient pattern was conserved during spring for fucoxanthin, but the increasing factor between winter and spring stocks was much less than for BSi and ranged from 2 to 3. Similarly to BSi, the highest fucoxanthin stock was observed in A1 in the spring (15 mg m^{-2}).

[20] The seasonal evolution of the integrated Chl *a* stocks were quite distinct from the diatom components (BSi and fucoxanthin) evolution, with already elevated stocks during winter period ($38\text{--}64 \text{ mg m}^{-2}$), followed by a very slight increase during spring. The integrated Chl *a* stocks increased by only 3% in A2 and 21% in C4 between P1 and P2. Surprisingly, Chl *a* stocks decreased by 24% in A1 during spring, contrary to the strong BSi and fucoxanthin increase at this site. In the late summer, the BSi and pigments stocks were lower than winter values at all sites, and were homogeneously distributed from north to south, with a slightly higher remaining fucoxanthin and Chl *a* stocks in A1 than at the other sites.

[21] Integrated primary production (ΣPP over 80 m) distribution was closer to Chl *a* than to the siliceous components distribution. The south-north increasing gradient was clear during winter, with a 35% increase of the ΣPP between A2 and A1. In the spring however, the ΣPP distribution differed from that of other parameters. The strongest increase in ΣPP was observed in A2 (+32%) while the smallest increase was measured in A1 (<2%), where the ΣPP was already high during winter ($62.1 \text{ mmol C m}^{-2} \text{ d}^{-1}$). Furthermore, the ΣPP decreased in C4 from

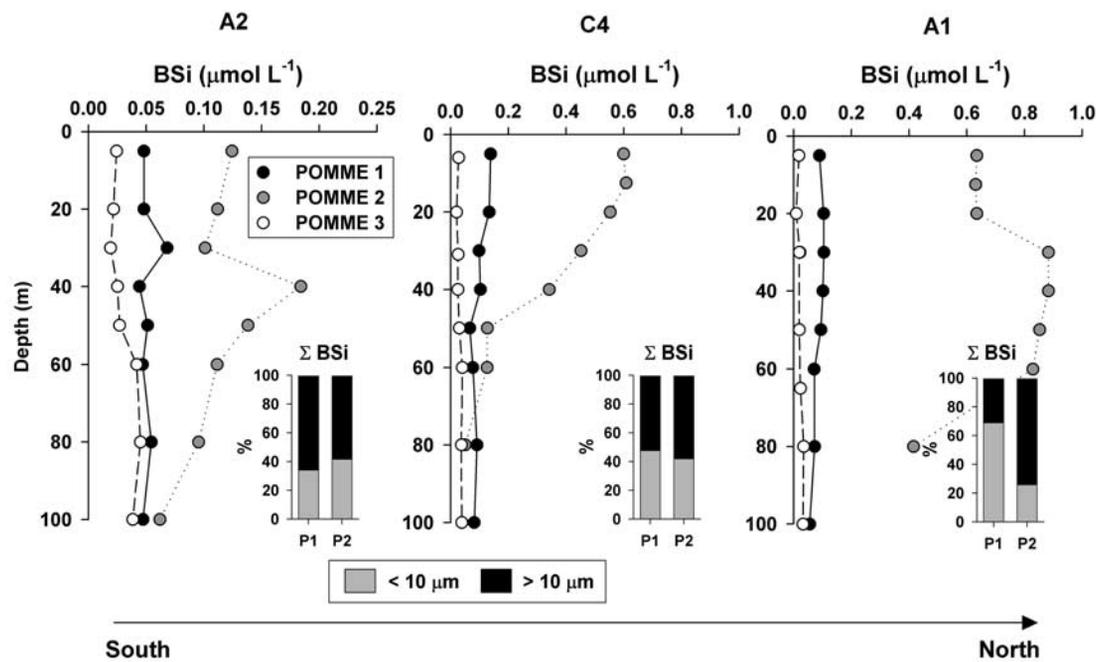


Figure 4. Biogenic silica (*BSi*) profiles for the 0–100 m layer for the three cruises at the southern anticyclonic eddy A2, the cyclonic eddy C4, and the northern anticyclonic eddy A1. The histograms present the percentage of small and large diatoms (<10 and >10 μm) to integrated *BSi* (ΣBSi) for POMME 1 (P1) and POMME 2 (P2).

53.9 $\text{mmol C m}^{-2} \text{d}^{-1}$ in the winter to 38.9 $\text{mmol C m}^{-2} \text{d}^{-1}$ in the spring (-28%). The lowest ΣPP values were measured in the late summer and ranged from 12.9 to 18.9 $\text{mmol C m}^{-2} \text{d}^{-1}$ from south to north.

3.4. Si Uptake Rates and Uptake Kinetics (V_{max} , K_S)

[22] The seasonal evolution of silicic acid uptake rates for A2, C4 and A1 is presented in Figure 6. In the winter, uptake rates were low in the first 50 m and reached a maximum of 0.04 $\mu\text{mol Si L}^{-1} \text{d}^{-1}$ in A1 at 20 m. In the spring, Si uptake rates increased significantly in the surface layer, but remained low at both A2 and C4 sites, contrary to *BSi* concentrations, which increased sharply in C4 (Figure 4). The increase in Si uptake rates followed a south-north gradient, and the highest values were measured in A1 at the surface (0.33 $\mu\text{mol Si L}^{-1} \text{d}^{-1}$) while a secondary peak was observed at 30 m (0.25 $\mu\text{mol Si L}^{-1} \text{d}^{-1}$). Si uptake rates were the lowest ever measured in any open ocean system during late summer cruise, with uptake rates below 0.001 $\mu\text{mol Si L}^{-1} \text{d}^{-1}$ at all sites over the entire 0–100 m layer.

[23] Size-fractionated integrated Si uptake rates over the 0–80 m layer showed a strong northward increase (from 34 to 81%) of the contribution of small diatoms (<10 μm) to the production fluxes during winter (Figure 6). In the spring, the contribution of small diatoms was higher in A2 and C4 than in the winter, while a strong shift toward the dominance of large diatoms was observed in the A1, where diatoms >10 μm contributed to 91% of the Σ Si uptake flux.

[24] Si uptake kinetic experiments were also conducted at two depths at each study site during P1 and P2. Out of the 16 kinetic curves obtained, 12 could be fitted by a Michael-

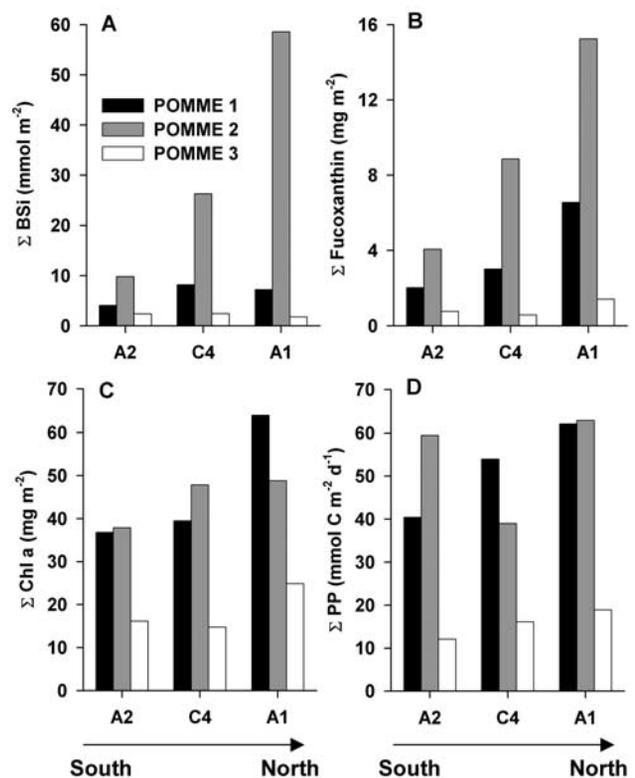


Figure 5. Integrated stocks (over 80 m) of (a) *BSi* (mmol m^{-2}), (b) fucoxanthin (mg m^{-2}), (c) *Chl a* (mg m^{-2}), and (d) integrated fluxes of primary production (PP) ($\text{mmol C m}^{-2} \text{d}^{-1}$) at the three revisited sites in the winter (P1), spring (P2), and late summer (P3).

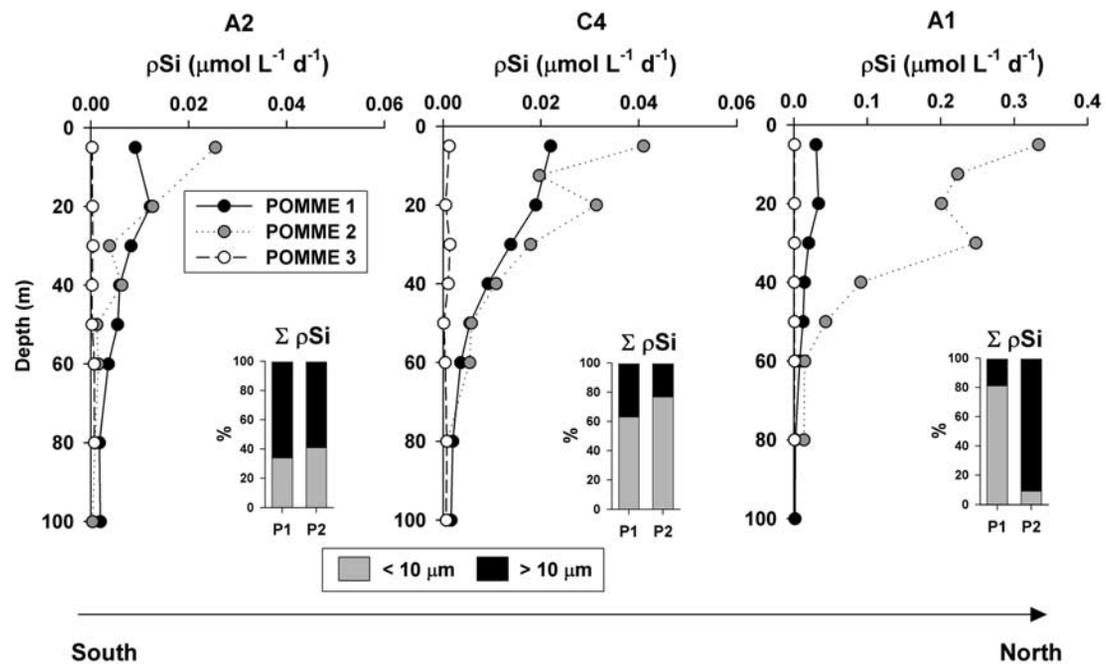


Figure 6. Si uptake rates profiles for the 0–100 m layer for the three cruises at the southern anticyclonic eddy A2, the cyclonic eddy C4, and the northern anticyclonic eddy A1. The histograms present the percentage contribution of small and large diatoms (<10 and >10 μm) to integrated Si uptake rates ($\Sigma\rho\text{Si}$ in $\text{mmol Si m}^{-2} \text{d}^{-1}$) for P1 and P2.

lis-Menten curve (hyperbolic fit), 3 presented a linear response to increasing silicic acid concentrations, and 1 already reached plateau values for the minimal Si addition, thus not allowing to calculate V_{max} or K_S parameters. Examples of the two main types of kinetic curves obtained (hyperbolic and linear) are presented in Figure 7. The calculated V_{max} and K_S values are summarized in Table 2 and compared to measured in situ $V\text{Si}$ and silicic acid concentrations. During P1, Si half-saturation constants were particularly high, ranging from 1.77 to 5.92 μM . The potential maximum specific Si uptake rates (V_{max}) ranged from 0.19 to 0.59 d^{-1} , with an average value of 0.43 d^{-1} (Table 2). In this season, K_S Si were always superior to ambient silicic acid concentrations values while measured

$V\text{Si}$ values represented between 22 and 65% of the calculated V_{max} . During P2, K_S Si values decreased significantly ($P < 0.0005$, $df = 7$, t-test) at all sites and were composed between 0.72 and 1.26 μM , which was slightly inferior to ambient silicic acid concentrations. The average V_{max} value was 0.34 d^{-1} , but the decrease between the winter and spring mean values was not significant ($P > 0.05$). The measured $V\text{Si}$ values were again lower than the calculated maximal potential uptake rates and represented 11–46% of the V_{max} values. The V_{max}/K_S ratio, proportional to the initial slope of the curve, is often considered to be a better nutrient affinity indicator than either V_{max} or K_S since both parameters can constrain the ability to take up nutrients at low concentrations [Healey, 1980; Franck et al., 2003]. The

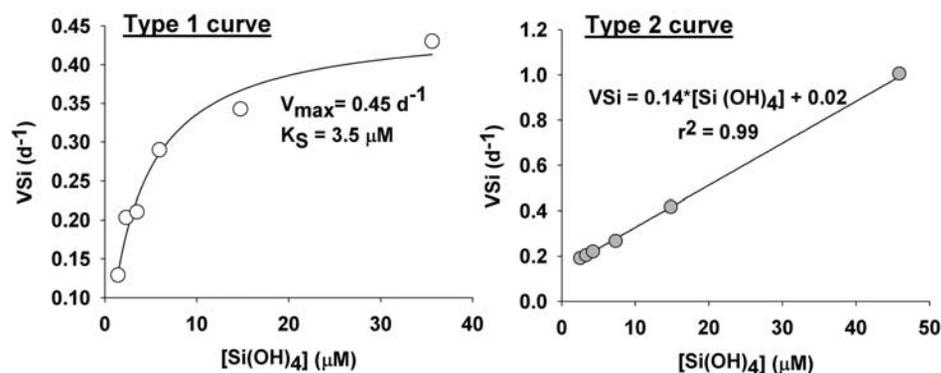


Figure 7. Examples of the two hyperbolic and linear types of kinetic Si uptake curves obtained during POMME 1 and 2. $V\text{Si}$, specific Si uptake rates (in d^{-1}).

Table 2. Si Uptake Kinetics During the Winter and Spring Season

	K_S Si ^a		[Si(OH) ₄] ^a		V_{max} ^b		VSi ^b		VSi/V_{max} ^b %		V_{max}/K_S	
	20 m	40 m	20 m	40 m	20 m	40 m	20 m	40 m	20 m	40 m	20 m	40 m
	<i>P1</i>											
A2	3.50	1.77	1.58	1.72	0.45	0.21	0.25	0.13	55	65	0.13	0.12
FZ	4.73	5.92	1.87	1.92	0.56	0.59	0.19	0.13	34	22	0.12	0.10
C4	4.10	4.50	2.37	2.12	0.41	0.19	0.14	0.09	34	47	0.10	0.04
A1	4.62	lin	2.40	2.50	0.61	unsat	0.32	0.14	52	...	0.13	...
	<i>P2</i>											
A2	lin	lin	1.53	1.47	unsat	unsat	0.20	0.03
C4	0.94	0.84	1.53	2.47	0.54	0.14	0.06	0.03	11	21	0.57	0.17
SP	1.26	0.85	1.36	1.32	0.55	0.24	0.23	0.11	42	46	0.44	0.28
A1	nd	0.72	1.29	1.03	nd	0.25	0.32	0.10	...	40	...	0.35

^aComparison of half-saturation constants for silicic acid (K_S Si in μM) with in situ silicic acid concentrations ([Si(OH)₄] in μM) for two depths at each study site during POMME 1 and 2 (P1 and P2): lin, linear uptake kinetics (type 2 curve in Figure 8); nd, not determined, no curve could be fitted to the data points.

^bComparison of maximum specific Si uptake rates (V_{max} in d^{-1}) with in situ specific Si uptake rates (VSi in d^{-1}) for two depths at each study site during POMME 1 and 2: unsat, linear response (type 2 curve in Figure 8), no plateau reached; nd, not determined.

V_{max}/K_S ratios appeared significantly increased in spring compared to winter values, with an increase on average by a factor 4–5 at 20 m and by a factor 3 at 40 m.

3.5. Diatom Contribution to Primary Production

[25] Diatom contribution to primary production was estimated using an average Si/C ratio to convert integrated Si uptake rates into C equivalents. A Si/C ratio of 0.09 was used for the small diatom fraction (<10 μm) and a Si/C ratio of 0.13 was used for the large diatoms [Brzezinski, 1985]. The converted C uptake rates were compared to the actual in situ ΣPP measurements and the contribution of diatoms to total phytoplankton ΣPP was mapped for each cruise (Figure 8). As indicated by other siliceous biomass components (BSi, fucoxanthin, ρSi), the contribution of diatoms to ΣPP increased from south to north at all seasons. This contribution increased from 12% in A2 to 25% in A1 during winter. In the spring, diatoms contribution to ΣPP remained unchanged in the A2, but increased drastically northward

and accounted for 100% of the ΣPP in A1. During late summer, this contribution decreased sharply and was insignificant over the entire study area (2–4%).

3.6. Si Export Fluxes

[26] BSi daily export fluxes were measured in drifting sediment traps at 200 m and 400 m at each study site (Figure 9). BSi fluxes were corrected for dissolution in the collecting cups. The percentage of BSi dissolution ranged between 0 and 50% throughout the study. The measured dissolution rates were moderate during winter (18%), increased during spring (34%) and were minimal during late summer (9%). In terms of absolute quantity, the corrected BSi fluxes were reevaluated on average by 22% compared to the uncorrected values.

[27] During both the winter and spring, BSi daily fluxes increased from south to north, following the pattern observed for biomass and production rates. In the winter, BSi fluxes at 200 m increased from 0.2 to 29.0 $\mu\text{mol Si}$

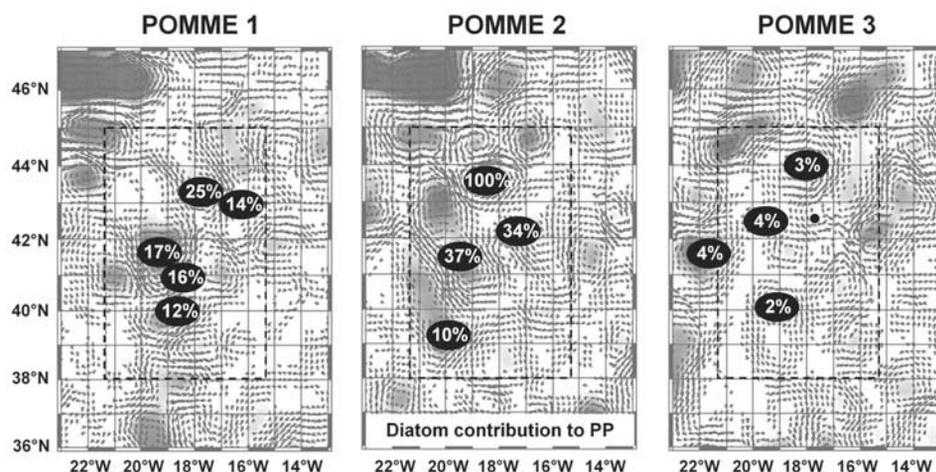


Figure 8. Diatoms contribution to integrated primary production (0–100 m) at the different study sites for each cruise. Si/C ratios of 0.09–0.13 were used to convert Si production into C equivalent.

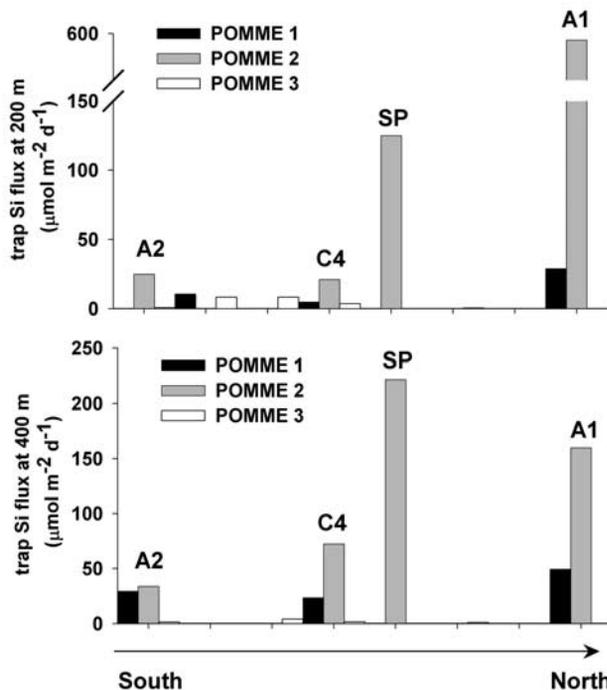


Figure 9. Biogenic silica fluxes ($\text{mmol m}^{-2} \text{d}^{-1}$) in sediment traps at 200 and 400 m along a south-north gradient for all three seasons.

$\text{m}^{-2} \text{d}^{-1}$ from A2 to A1. *BSi* fluxes increased significantly during spring, and ranged from 24.8 to 586.0 $\mu\text{mol Si m}^{-2} \text{d}^{-1}$ from south to north. During P1 and P2, fluxes at 400 m were always superior to the fluxes measured at 200 m, except in A1, where the flux decreased by a factor 3.6. During spring, the maximum Si flux at 400 m was recorded at 42°N (221.8 $\mu\text{mol Si m}^{-2} \text{d}^{-1}$) and not in A1, where the upper trap collected the highest flux. In the late summer, fluxes were negligible at both depths (mean Si flux: 2.7 $\mu\text{mol Si m}^{-2} \text{d}^{-1}$).

4. Discussion

4.1. Onset of the Bloom: The South-North Gradient Compared to Mesoscale Features

[28] The first phase of the bloom occurs after the alleviation of light limitation in the North Atlantic, following the spring stratification of the surface layer [Sverdrup, 1953]. Heat fluxes toward the ocean became positive at the end of February [Caniaux et al., 2005a, 2005b]. The CARIOCA buoys dispatched throughout the study area indicated the start of a biological uptake of CO_2 by 1 April, in parallel with an increase in fluorescence [Mémery et al., 2005]. This discrepancy may be explained by the frequency and intensity of wind stress in this region, likely to delay the onset of stratification in the surface layer [Stramska et al., 1995]. Satellite derived zonal integrations of *Chl a* concentrations indicated a northward progression of the spring bloom between April and May, when the peak *Chl a* level progressed from 40°–41°N to 45°N [Lévy et al., 2005]. Each leg of the POMME cruises were carried out over 25 days and 4° of latitude, meaning that the synopticity of these results is not likely. Thus the south-north gradient

observed for most of our data results partly from (1) a spatial variability, induced by the differences in available nutrient stocks created by the northward increase of the winter mixed layer depth (MLD) and (2) a temporal variability, as the bloom progresses northward following increased heat fluxes and the propagation of the thermal stratification. The interpretation of the data from the second legs is thus relatively complex, but the diatoms distribution appeared more influenced by the latitudinal progression of the bloom rather than by the influence of the different mesoscale features (anticyclonic and cyclonic eddies). Indeed, the siliceous biomass, Si uptake rates and diatoms contribution to primary production increased regularly from south to north (Figures 4, 5, 6, and 8).

4.2. Diatom Seasonal Dynamics in the Eastern North Atlantic

4.2.1. Diatom Distribution

[29] Diatom biomass indicators as well as Si production fluxes were poorly correlated to any other parameter such as nutrient stocks and primary production. In an attempt to reconstruct a temporal evolution of the diatom bloom, integrated *BSi*, fucoxanthin and ρSi (0–80 m) were represented following a decreasing integrated (0–80 m) silicic acid gradient by combining data for all sites and all seasons (Figure 10). According to Brown et al. [2003], Si acid uptake rates could be predicted from the average photic silicic acid concentrations in the North Atlantic. This relationship was not evident in our study, whether we considered the mean photic zone average concentrations (data not shown) or the 0–80 m integrated stocks of silicic acid (Figure 10). Integrated *BSi*, fucoxanthin and Si uptake fluxes were highly correlated throughout the survey, but did not show any correlation with integrated silicic acid, nitrate or *Chl a*. However, both nitrate and silicic acid integrated stocks showed a good correlation with integrated *Chl a*. Thus a more detailed study of nutrient distribution pattern was necessary to investigate the main factor controlling the evolution of the diatom bloom.

4.2.2. Limiting Factors: Evidence for Si Limitation

[30] In the spring, the North Atlantic production regime is mainly controlled by the availability of nutrients once the surface layer stratification has occurred [Longhurst et al., 1995]. In the North Atlantic, the potential limiting role of silicic acid of the diatom spring bloom has been pointed out through field studies and modeling work [Lochte et al., 1993; Sieracki et al., 1993; Passow and Peinert, 1993; Taylor et al., 1993; Savidge et al., 1995; Bury et al., 2001].

[31] During winter, silicic acid availability was already 2–3 μM lower than nitrate and integrated stocks were on average twice as low (135–200 mmol Si m^{-2} compared to 280–415 mmol N m^{-2}). This deficit in Si was already observed at 1000 m in deep waters, where nitrate concentrations were 5–7 μM higher than silicic acid values (deep profiles not shown), implying a smaller initial input of Si compared to N to the surface layer through winter vertical mixing. The low Si:N ratios in intermediate waters are explained by a slower dissolution rate of *BSi* during particulate matter sinking compared to N regeneration. Along the conveyor belt circulation through the world oceans, waters below the euphotic zone tend to accumulate Si which is less readily recycled in the water column than N.

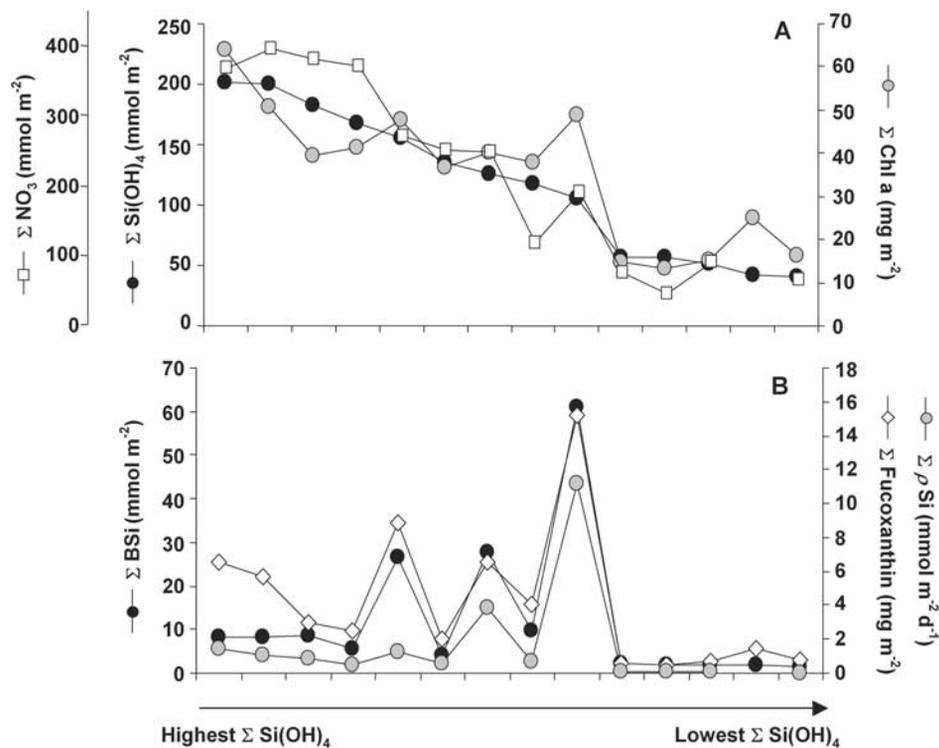


Figure 10. Combining the results from all three seasons, the data for each site presented in order of decreasing integrated (0–80 m) silicic acid gradient (mmol m^{-2}) following the x axis. (a) Corresponding integrated nitrate (mmol m^{-2}) and Chl a (mg m^{-2}) and (b) BSi (mmol m^{-2}), fucoxanthin (mg m^{-2}), and Si uptake rates ($\text{mmol Si m}^{-2} \text{d}^{-1}$) values (0–80 m) for each site plotted according to the same decreasing silicic acid gradient.

A steady increase of the Si:N nutrient ratio at 250 m has thus been observed between the starting point of the conveyor belt in the Atlantic Ocean with Si:N ratios close to 0.5 and the North Pacific Ocean which receives older deep waters with Si:N ratios >1.5 [Ragueneau *et al.*, 2000]. The deficit in Si in the North Atlantic seems to be a recurrent feature and was also observed in the western basin in the Sargasso Sea by Brzezinski and Nelson [1996] who reported a mean Si:N ratio at the basis of the euphotic layer as low as 0.25. During the North Atlantic Bloom Experiment (NABE) program, Sieracki *et al.* [1993] indicated a faster depletion of silicic acid relative to nitrate following the onset of the spring bloom at 46°N and 18°W , just north of the POMME study area. These authors measured silicic acid concentrations as low as $0.2 \mu\text{M}$ in May, which is extremely low for an open ocean regime. Comparable values were also recorded in the subtropical region of the North Atlantic (33°N – 21°W) by Passow and Peinert [1993], who documented concentrations $<0.2 \mu\text{M}$ over the entire 0–200 m layer, while the depth of the nitracline oscillated between 60 and 80 m. During the POMME cruises, such low silicic acid concentrations were only observed during late summer survey, when dissolved Si varied between 0.2 and $0.6 \mu\text{M}$ in the 0–50 m layer. However, it is likely that Si was already depleted in the late spring, as both the CARIOCA buoys deployed in the study area and the Sea-viewing Wide Field-of-view Sensor (SeaWiFS) images indicated that the peak of the spring bloom occurred in April in the south and in May in the

north of the POMME area, the latter being only partially sampled by the end of the second leg of P2.

[32] In order to determine the potential limiting factor of the diatoms spring bloom, in the sense of Liebig's law of minimum, we examined the integrated nutrient ratios in the surface layer compared to the empirical ratio Si:N:P = 16:16:1 [Brzezinski, 1985; Redfield *et al.*, 1963], which gives an estimate of their nutrient requirement. Integrated Si:N and N:P ratios were calculated over the 0–80 m layer for all sites. In the winter, the Σ Si:N ratios were systematically <0.5 and increased slightly during spring (Figure 11a). The highest value was observed in A2 (0.74) where nitrate was depleted more rapidly. In the late summer, silicic acid availability compared to nitrate increased in C4 (Σ Si:N = 1.12) and particularly in A1 (Σ Si:N = 5.30). At the other sites, the ratio remains <1 , indicating a concomitant depletion of silicic acid and nitrate. The Σ N:P ratios indicated a slight P deficit in the winter period (Figure 11b), with ratios >16 in A2, C4 and A1 eddies (17.3–21.1). In the spring and late summer, the Σ N:P ratio indicated a constant deficit of N compared to P with much lower values (7.5–12.8).

[33] All data points collected during the POMME cruises for which nitrate, silicic acid and phosphate concentrations were measured concurrently between 0 and 80 m were summarized in Figure 12, where the Si:N = 1, Si:P = 16 and N:P = 16 lines delimit six areas, each characterizing a different hierarchy for the potential limiting factors. A seasonal shift of the potential limiting nutrient is clear: in

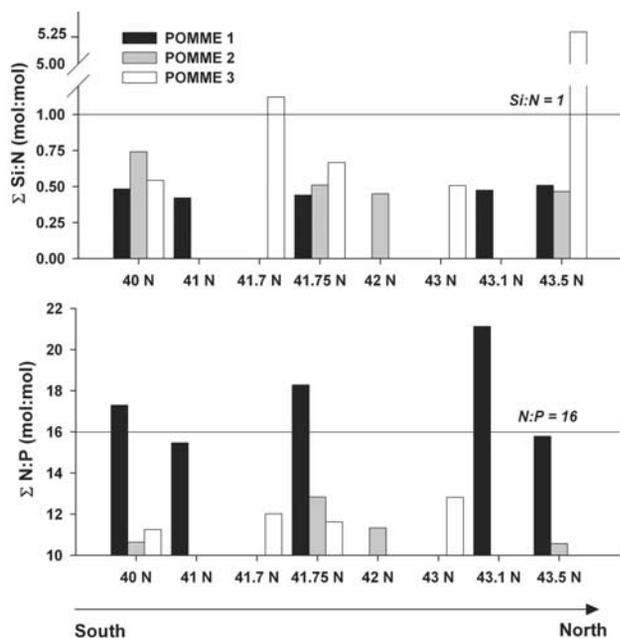


Figure 11. Integrated silicic acid:nitrate ratios (mol:mol) and nitrate:phosphate ratios (mol:mol) along a south-north gradient for the three seasons.

the winter, Si appears to be the first potential limiting nutrient, then followed by P and N in decreasing importance. In the spring, Si remained the first potential limiting nutrient while N was the secondary limiting nutrient followed by P. In the late summer, the nutrient status changed drastically toward a strong primary limitation by N.

[34] While the comparison of ambient nutrients with the empirical Si:N:P ratio gives an indication on the hierarchy of the potential limiting nutrients, this approach is not sufficient in itself to determine the degree of limitation by one element. Nutritional limitation depends on an equilibrated supply of Si:N:P but also on the physiological status of phytoplankton, which can present various degrees of adaptation to low nutrient levels, on the specific community structure as well as on regeneration rates [Del Amo *et al.*, 1997].

[35] Another way to assess the degree of limitation is to characterize the uptake kinetics of the considered element and to compare half-saturation constants (K_S) with in situ concentrations. The Si uptake kinetic experiments conducted during winter and spring allowed us to give the first estimates for K_S Si values in the northeast Atlantic (Table 2). These results indicated relatively high values of K_S Si in the winter, with a mean $K_S = 4.16 \mu\text{M}$, in the high range of the values given in the literature ($0.4\text{--}5.0 \mu\text{M}$) by Nelson and Tréguer [1992].

[36] The poor kinetic efficiency of diatoms may have been explained by light limitation at this season, but the thermal stratification already allowed a nonsiliceous phytoplankton to develop, as evidenced by the elevated $\Sigma \text{Chl } a$

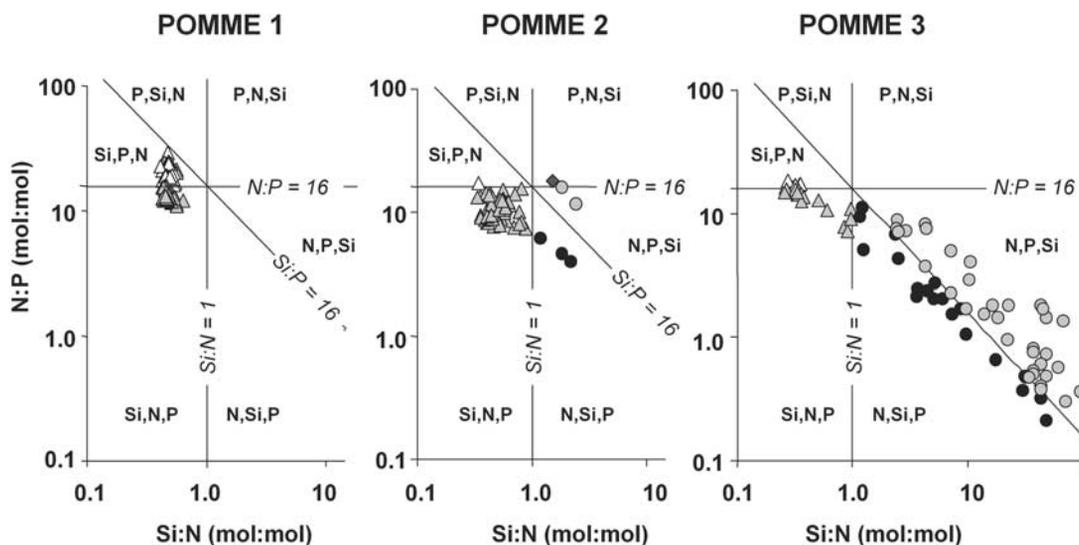


Figure 12. Si:N:P molar ratios of all nutrient data collected during each POMME cruise for the 0–80 m layer. Each area of the graph is delimited by straight lines representing the empirical Si:N:P = 16:16:1 ratios [Brzezinski, 1985; Redfield *et al.*, 1963]. Data with Si:N values <1 show potential Si limitation, data with N:P values <16 show potential N limitation, etc. A subsequent hierarchy of the order of importance of the potential limiting nutrient (PLN) can thus be attributed to each of the six areas delimited by the Si:N:P ratios. Different symbols were used for each area: open triangles, Si first potential limiting nutrient, followed by P and N in decreasing order of importance; shaded triangles, Si first potential limiting nutrient, followed by N and P in decreasing order of importance; solid circles, N first potential limiting nutrient, followed by Si and P; shaded circles, N first potential limiting nutrient, followed by P and Si; solid diamonds, P first potential limiting nutrient, followed by N and Si.

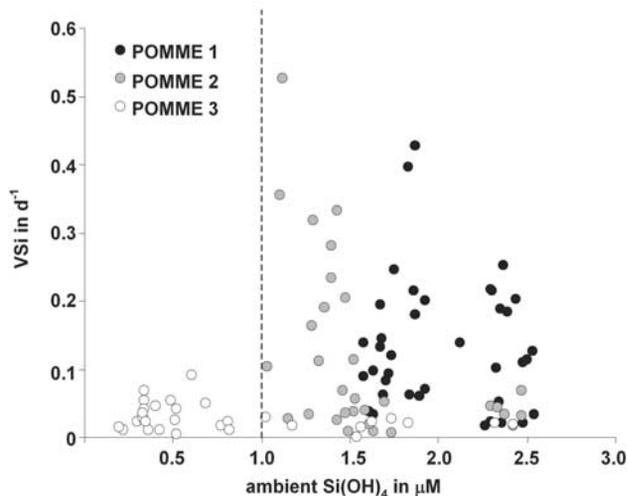


Figure 13. Specific Si uptake rates (d^{-1}) between 0 and 80 m as a function of in situ silicic acid concentrations (in μM). The straight line delimits conditions for which silicic acid is $<1 \mu\text{M}$.

and Σ PP in the winter, while Σ BSi, Σ fucoxanthin and Si uptake rates were still very low (Figures 5 and 6). These results suggest that the alleviation of light limitation was already occurring in the winter, allowing the early onset of the annual phytoplankton bloom, but that diatoms growth was subject to strong Si limitation, as indicated by the K_S , V_{max} and V_{max}/K_S values as well as by the deficit in Si in the surface layer compared to other nutrients.

[37] In the spring, the lower mean K_S value ($0.92 \mu\text{M}$) and the increase in the affinity ratio (V_{max}/K_S) compared to the winter values suggest a better adaptation of diatoms to the low ambient Si concentrations. Nonetheless, the increase of Si uptake rates remained heterogeneous over the study area, with very low ρSi in the southern/central region (A2 and C4) and a large increase reflecting a bloom situation in the north (A1). While VSi still appeared strongly limited in that season, the increase of the $V\text{Si}/V_{\text{max}}$ ratio was higher in A1 (40%) than in C4 (11–21%), confirming diatoms increased growth rate and Si uptake efficiency at the northern site. It is likely that diatoms also encounter N limitation in the southern region, where nitrate decreased sharply between the winter and spring and where the Σ Si:N ratios are highest (Figure 11). Nitrate concentrations were $<0.5 \mu\text{M}$ in the surface layer, which is potentially limiting if one considers the average K_S NO_3^- given in the literature of $2 \mu\text{M}$ [Del Amo et al., 1997]. In the North Atlantic, Savidge et al. [1995] attributed the decrease of diatoms growth rates to the diminution of silicic acid concentrations to values $<1 \mu\text{M}$ and suggested simultaneous limitation of diatoms by both Si ($<0.8 \mu\text{M}$) and N ($<0.5 \mu\text{M}$). During NABE, Sieracki et al. [1993] also attributed the decline of a spring *Pseudo-Nitzschia* bloom to the low availability of silicic acid.

[38] According to Nelson and Dortch [1996], the conditions least favorable to a diatom bloom are reached whenever silicic acid concentrations fall below $1.5 \mu\text{M}$ in parallel to Si:N ratios <1 . Despite these nonfavorable conditions in the northern region ($\text{Si(OH)}_4 < 1.3 \mu\text{M}$;

$\text{Si:N} < 0.5$), a *Pseudo-Nitzschia* bloom was observed in the A1 eddy during spring. This atypical bloom may result from a higher hivernal silicic acid stock in the surface layer compared to the other sites, as well as from a better adaptation of diatoms to the ambient silicic acid availability, as evidenced by the lowest K_S value measured during spring ($0.72 \mu\text{M}$). Furthermore, diatoms may adapt to low Si conditions by reducing the thickness of their frustules, since Si half-saturation constants for growth (K_{μ}) are often much lower than half-saturation constants for uptake rates (K_S) [Paasche, 1973; Nelson et al., 1976].

[39] The uptake of Si by diatoms has been demonstrated to follow a Michaelis-Menten curve with increasing substrate concentrations (type I curve (Figure 7)) [Paasche, 1980]. However, several studies have documented linear responses of uptake rates with increasing silicic acid concentrations comparable to the type II curve also observed in our study (Figure 7). Similar uptake kinetics were documented in the tropical Atlantic waters [Brzezinski and Nelson, 1996], in the central North Pacific [Brzezinski et al., 1998] and in the northeast Atlantic [Brown et al., 2003]. In this last study, the authors discuss the possibility that the linear fit represents the lower section of the hyperbolic curve. This is very unlikely in our study, since Si(OH)_4 additions ranged from 0.5 to $50 \mu\text{M}$, while K_S values for oceanic systems (except in the Southern Ocean) are usually constrained between 0.4 and $5 \mu\text{M}$ [Nelson and Tréguer, 1992]. In A1 during spring, the surface diatom community furthermore exhibited saturated uptake rates already at low silicic acid concentrations and K_S and V_{max} could not be determined. These different types of responses may reflect distinct Si uptake models such as surge uptake, externally or internally controlled uptake [Martin-Jézéquel et al., 2000], which we hypothesize could depend on the relative external and internal Si pool as well as on the physiological state of the cells, and to some degree on a specific genetic variability. Clearly further studies aiming at identifying and characterizing the chain of enzymes and transporters involved in Si uptake are needed to understand the linear responses of Si uptake kinetics now documented for several oceanic systems.

[40] In situ measured VSi values did not show any linear relationship with ambient silicic acid concentrations (Figure 13) contrary to what was evidenced by Brzezinski et al. [1998] in the North Pacific. In this study, authors found that specific uptake rates were close to zero whenever silicic acid concentrations were $<1 \mu\text{M}$. Despite the lack of clear relationship between VSi and silicic acid during the POMME experiment, we still observed that VSi were always very low (0.1 d^{-1}) below $1 \mu\text{M}$ of Si, indicating a strong limitation of uptake rates below this threshold. However, the sharp diminution of VSi values was observed during late summer and may also be the consequence of the extreme oligotrophic conditions observed, where nitrate and phosphate levels were low enough to initiate the decline of the microphytoplankton population. In the late summer, the increase of silicic acid concentration up to $2.5 \mu\text{M}$ did not result in any further increase of VSi values (Figure 13), confirming the primary limiting role of nitrate and phosphate at this season. Similarly, the absence of linear relationship between VSi and dissolved Si in the winter and spring suggest the occurrence of simultaneous limitation by more

Table 3. Comparative Table of Integrated Si Uptake Fluxes for the Entire POMME Area and Other Systems of the North Atlantic

Study Site	Latitude, °N	Month	$\Sigma\rho\text{Si}$, $\text{mmol m}^{-2} \text{d}^{-1}$		References
			Range	Mean	
North Atlantic					
POMME 1	40–43.5	March	0.55–1.45	0.90	
POMME 2	40–43.5	April	0.67–11.20	4.25	this study
POMME 3	40–43.5	Sept.–Oct.	0.04–0.08	0.07	
Faeroes-Iceland-Scotland (FISHES)	56–64	May	6–167	35.5	<i>Brown et al.</i> [2003]
North Atlantic (PAP site)	48.8		0.5–1.3	0.9	<i>Ragueneau et al.</i> [1997]
Sargasso Sea	31.5	yearly	0.2–1.6	0.5	<i>Brzezinski and Nelson</i> [1995]; <i>Brzezinski and Kosman</i> [1997]
		April	4.4–9.8	7.1	
Gulf Stream		June	5.3–11.7	7.3	<i>Brzezinski and Nelson</i> [1989]
		Sept.–Oct.	2.0–5.0	3.4	

than one element. Shipboard enrichment experiments conducted during the POMME experiment suggested that Si-Fe colimitation processes were also likely to limit diatom growth in this area [Blain *et al.*, 2005], which is not usually considered to be an Fe-limited region due to the proximity of the continental shelf. However, low in situ Fe concentrations measured in the study area and a strong response to additions of both Fe and Saharan dusts on diatoms growth [Blain *et al.*, 2005] contradicts the precedent assumption that the North Atlantic does not encounter Fe limitation.

4.2.3. Seasonal Si Production and Export Budgets

[41] Mean integrated Si production fluxes were calculated for each season in the POMME study area and compared to other production values measured in the North Atlantic (Table 3). There exists only a few studies documenting Si uptake rates in the North Atlantic and only two were carried out in the northeastern Atlantic. Integrated Si uptake rates measured north of our study area at the PAP site (48.8°N) were very similar to our Si production budget calculated for the winter period with the same $\Sigma\rho\text{Si}$ average of 0.9 $\text{mmol Si m}^{-2} \text{d}^{-1}$ [Ragueneau *et al.*, 1997]. The range values of winter Si production fluxes was also very close to that found in the Sargasso Sea [Brzezinski and Nelson, 1995; Brzezinski and Kosman, 1997]. On the other hand, a study conducted between Iceland and Scotland between 56 and 64°N documented the highest production rates measured in the North Atlantic Ocean (maximum of 167 $\text{mmol Si m}^{-2} \text{d}^{-1}$) [Brown *et al.*, 2003], one to two orders of magnitude higher than the rates measured in the present study. During the productive period, the range of $\Sigma\rho\text{Si}$ values during POMME was in the same order of magnitude than those

found in warm eddies of the Gulf Stream [Brzezinski and Nelson, 1989]. However, the average $\Sigma\rho\text{Si}$ for that period was 40% lower than the production fluxes measured in April and June in the Gulf Stream. Despite the fact that the POMME study area is considered to be mesotrophic, with yet undepleted nutrients during spring bloom, Si production rates were extremely low and similar to the rates found in the oligotrophic gyres of the western Atlantic (Sargasso Sea and warm core eddies of the Gulf Stream), which are regions considered to encounter severe Si limitation throughout the year [Brzezinski and Nelson, 1996].

[42] The particulate silica export fluxes (BSi_{exp} flux) were compared to other studies also conducted in the North Atlantic at similar latitudes (Table 4). Very similar BSi_{exp} fluxes were obtained between the flux measured in drifting sediment traps deployed at each site at 400 m during the second legs and the average fluxes measured in moored sediment traps deployed over one year in the POMME study area [Mosseri *et al.*, 2005]. In most cases, BSi fluxes were measured at different depths (150–3000 m), but it has been shown that the BSi fluxes was conservative between the surface and deep traps in diverse oceanic systems, confirming the paradigm that diatoms are associated with rapid sedimentation events to depth [Ragueneau *et al.*, 2000]. Thus we expect our BSi_{exp} fluxes to be comparable with other studies. The average BSi_{exp} flux measured at 400 m during the POMME cruises in the spring (0.122 $\text{mmol Si m}^{-2} \text{d}^{-1}$) was very close to those found at the PAP site at 1000 m (0.160 $\text{mmol m}^{-2} \text{d}^{-1}$), north of our study area [Lampitt *et al.*, 2001]. The spring fluxes are also comparable to the annual average flux at the BATS site in the

Table 4. Comparative Table of Si Export Fluxes for the Entire POMME Area and Other Systems of the North Atlantic Between 150 m and 3000 m Depth

North Atlantic	Latitude, °N	BSi Flux, $\text{mmol m}^{-2} \text{d}^{-1}$					References
		150 m	200 m	400 m	1000 m	3000 m	
POMME 1	40–43.5		0.011	0.034			
POMME 2	40–43.5		0.189	0.122			this study
POMME 3	40–43.5		0.003	0.002			
POMME	40–43.5			0.039–0.147	0.090–0.111		<i>Mosseri et al.</i> [2005]
PAP	48.8				0.160	0.235	<i>Lampitt et al.</i> [2001]
PAP	48.8					0.139	<i>Ragueneau et al.</i> [2001]
PAP	48.8					0.096	
NABE prebloom	47.3				0.024	0.005	
NABE bloom	47.3				0.117	0.018	<i>Honjo and Manganini</i> [1993]
NABE postbloom	47.3				0.028	0.011	
JGOFS/BOFS	47.5					0.186	<i>Newton et al.</i> [1994]
BATS (Sargasso)	31.5	0.130	0.113	0.085		0.052	<i>Brzezinski and Nelson</i> [1995]

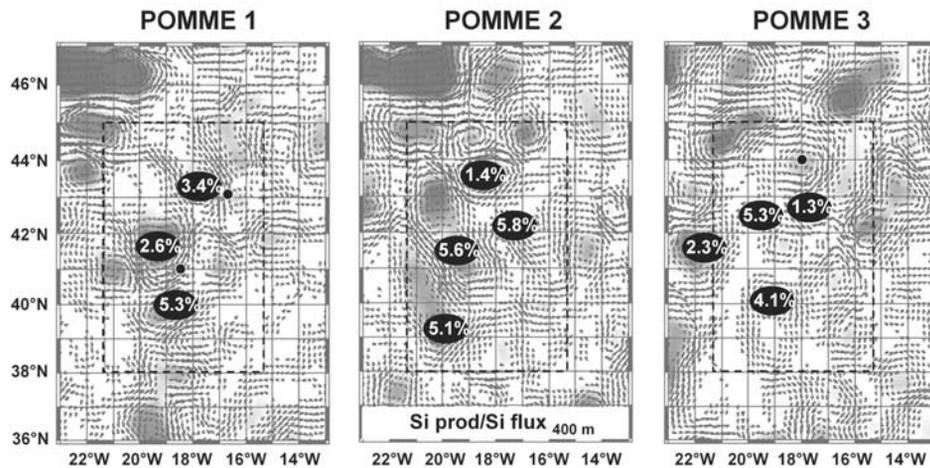


Figure 14. Percentage of the integrated Si production flux ($\Sigma\rho\text{Si}$) in the surface layer exported to 400 m depth.

Sargasso Sea [Newton *et al.*, 1994; Brzezinski and Nelson, 1995] and similar to the fluxes observed in the bloom period during NABE [Honjo and Manganini, 1993]. The fluxes reported for the PAP, BATS and BOFS study sites are however annual average fluxes, and may thus be considered higher than the BSi_{exp} fluxes observed in our study area, where fluxes were only elevated during the productive period, and very low during both winter and late summer. Similarly, during NABE, fluxes recorded during the non-productive period were much lower than the annual average at the PAP and BATS sites. The highest BSi_{exp} flux measured for the North Atlantic was recorded at 3000 m at the PAP site ($0.235 \text{ mmol Si m}^{-2} \text{ d}^{-1}$), which appears to be the site of important diatom sedimentation events. The differences in BSi_{exp} fluxes between PAP and POMME can be explained by the lesser intensity of the spring diatom bloom south of 45°N . SeaWiFS satellite images indicate that the POMME study area was located south of the maximal spring bloom extension, situated between 44° and 48°N (X. Pfeuty and M. Lévy, LODYC, personal communication, 2002). The north-south variability is thus important over a few degrees of latitude and diatom abundance appears to increase northward, with the increasing silicic acid availability. Koeve *et al.* [2002] also documented higher BSi concentrations at 47°N during spring, with values ranging between 74 and 258 mmol m^{-2} compared to $31 \pm 21 \text{ mmol m}^{-2}$ measured during P2 at 40° – 43.5°N .

[43] The percentage of biogenic silica produced in the surface layer and exported to 400 m was calculated ($\Sigma\rho\text{Si}/BSi_{\text{exp}400}$) and reported for each site in Figure 14. The percentage of Si exported to depth was higher during spring, with an average of 4.5% for all sites and was slightly lower during winter (3.8%) and late summer (3.2%). However, the fraction of exported biogenic silica did not vary significantly between the different season ($P > 0.001$, t-test) and the intrasite variability during each cruise was higher than the seasonal variability. The fraction of Si exported was less important in the A1 eddy despite the highest Si biomass and uptake rates observed. This may indicate either a more efficient recycling of the siliceous compartment in the surface layer or lateral advection processes as was suggested

for the annual BSi_{exp} flux at the northern site by Mosseri *et al.* [2005]. The fraction of Si exported to depth during the POMME cruises remains very low compared to the global average of 50% [Tréguer *et al.*, 1995] and to the average for the Southern Ocean of 40% [Ragueneau *et al.*, 2002] but is close to the values of 10% measured in the oligotrophic Equatorial Pacific [Blain *et al.*, 1999].

4.2.4. Annual Si Production and Export Budgets

[44] The average daily Si production and export rates were extrapolated to a yearly average over the POMME study area following the calculation described by Fernández *et al.* [2005]. On the basis of the nitrate stock, these authors estimated the length of the winter season to 49 days, the spring season to 59 days, and the rest of the year (257 days) was considered representative of the oligotrophic season. With a surface area for the POMME domain of $375\,000 \text{ km}^2$, the annual Si fluxes was extrapolated to $9.8 \times 10^9 \text{ mol Si yr}^{-1}$ while the export fluxes at 400 m was estimated at $2.9 \times 10^8 \text{ mol Si yr}^{-1}$. Thus a yearly average of 3% of the Si produced in the surface layer is exported at 400 m between 40° and 43.5°N in the North Atlantic.

5. Conclusion

[45] The results of the POMME cruises indicated an important seasonal variability of the siliceous compartment in the North Atlantic together with a spatial variability following a south-north gradient in parallel to the increasing stock of nutrients. The variability that could have been induced by the mesoscale hydrological features (anticyclonic and cyclonic eddies) was not apparent on particulate matter distribution and seemed masked by the south-north productivity gradient.

[46] In March, Chl *a* stocks were already elevated and equal to or slightly superior to the pigment values observed during April, indicating an early start of the spring bloom occurring before the complete stratification of the surface layer. Diatoms abundance was overall minor and their contribution to primary production was estimated to be $<25\%$ in the winter and $<4\%$ in the late summer. The low contribution of diatoms to the early bloom was attributed to

Si limitation over the entire study area. The dissolved Si:N ratio in the surface layer indicated a deficit of Si relative to N, and the elevated K_S Si values (1.77 to 5.92 μM) suggested that diatoms Si uptake rates were limited by silicic acid availability in the winter. In situ specific Si uptake rates (V_{Si}) were limited to between 22 to 65% of their potential maximum uptake rates (V_{max}).

[47] During spring, diatoms contribution to primary production increased northward, in parallel to the increasing silicic acid stock, ranging from 10% in the southern A2 eddy up to 100% in the northern A1 eddy, where a *Pseudo-Nitzschia* bloom was observed. Si uptake rates remained extremely low, except in A1, and seemed again limited by silicic acid availability. K_S Si values were lower (0.72–1.26 μM), reflecting a better adaptation of diatoms to decreasing concentrations over the course of the spring bloom, but V_{Si} were still limited from 11 to 46% of their V_{max} . In the southern part of the study area, an important drawdown of nitrate was likely to induce a simultaneous limitation of diatoms growth together with Si. In the late summer, nutrient depletion of the surface layer resulted in a strong decrease of the siliceous component, Chl *a* stocks and primary production, which was then dominated by the picophytoplanktonic fraction.

[48] The Si production budgets derived from in situ measurements were extremely low at all seasons (0.04–11.20 mmol Si $\text{m}^{-2} \text{d}^{-1}$) and closer to those established for other open ocean oligotrophic regimes such as the Sargasso Sea and the Gulf Stream rather than to those reported for productive areas. Several studies have already evidenced the poor uptake kinetic efficiency of diatoms in the North Atlantic and it has been hypothesized that the spring diatom bloom encounters chronic Si limitation [Brzezinski and Nelson, 1996; Brown et al., 2003]. The results of the POMME cruises tend to confirm these precedent studies, but simultaneous Si and N limitation were also likely to occur during spring, while nitrate depletion probably terminated the diatom bloom at the end of summer. Shipboard enrichment experiments described in a companion paper [Blain et al., 2005] also indicated that Fe may be limiting in this region.

[49] Si export fluxes to depth during the productive period were comparable to those measured in the same area on previous studies (PAP, NABE) but the average flux of the three seasons remained inferior to the annual fluxes measured between 47° and 48°N, where diatoms abundance was more important. The percentage of Si produced exported to 400 m varied between 1 and 6% over the POMME region and the intrasite variability was higher than the seasonal variability. Overall, Si export fluxes were very low compared to the global average of 50% for the open ocean.

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